

Differential gold endowment in
accretionary and dome-and-
basin Archean greenstone
belts

Enrichissement différentiel en
or des ceintures orogéniques
d'accrétion et dôme-et-bassin
d'âge Archéen

Bruno Lafrance

Congrès AEMQ XPLOR 2024



A new Canadian research initiative funded
by Canada First Research Excellence Fund.



What is Metal Earth?

- **METAL EARTH** is a collaborative research project focused on metal endowment in the Precambrian shield.
- It is led by the Mineral Exploration Research Centre (MERC) of Laurentian University.



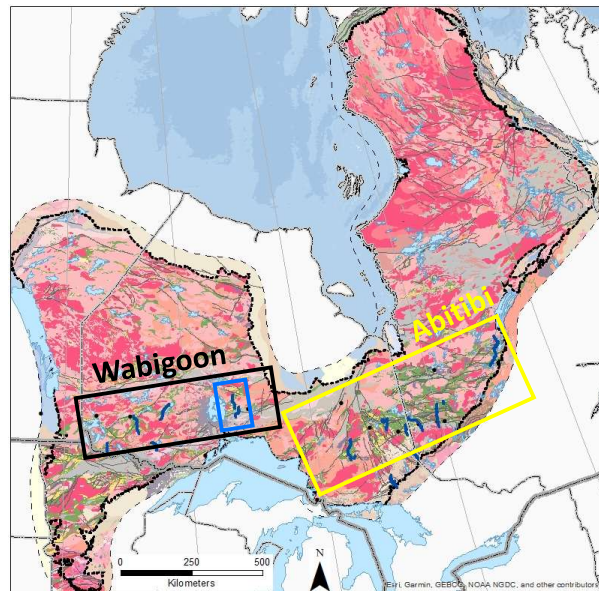
Metal Earth Four Main Components

- **Craton Scale Research:** investigate the architecture of greenstone belts and their link with surrounding granitoids during craton assembly and mineral district formation
- **Data Analytics Research:** Develop data integration, analysis and interpretive tools to predict metal endowment and guide exploration
- **Thematic Research:** to address specific processes or questions on metal endowment
- **Transect Scale Research:** resolve the lithospheric-crustal architecture and fluid pathways, providing a geological and geophysical framework to resolve the differential endowment of terranes and structures

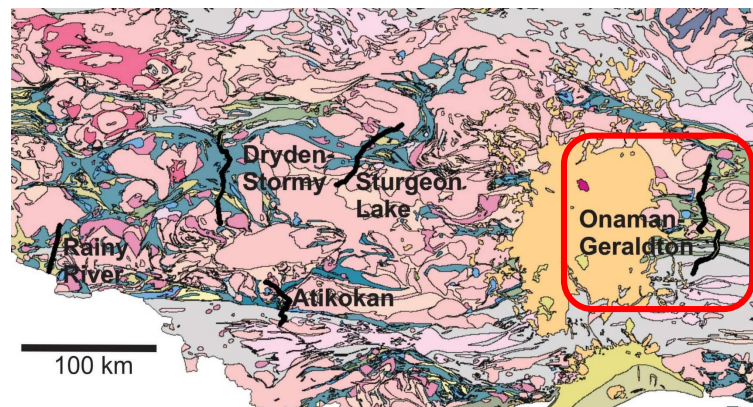
Transects were done across:

Endowed Abitibi Subprovince

Less endowed Wabigoon Subprovince



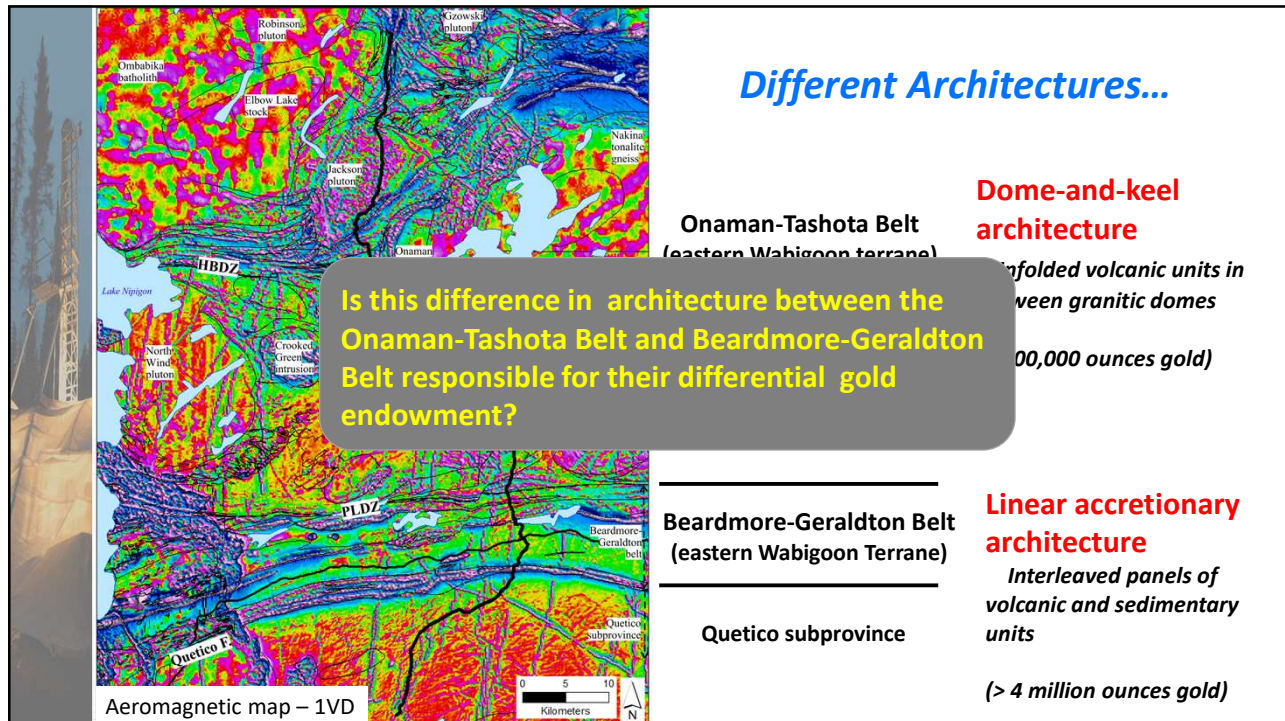
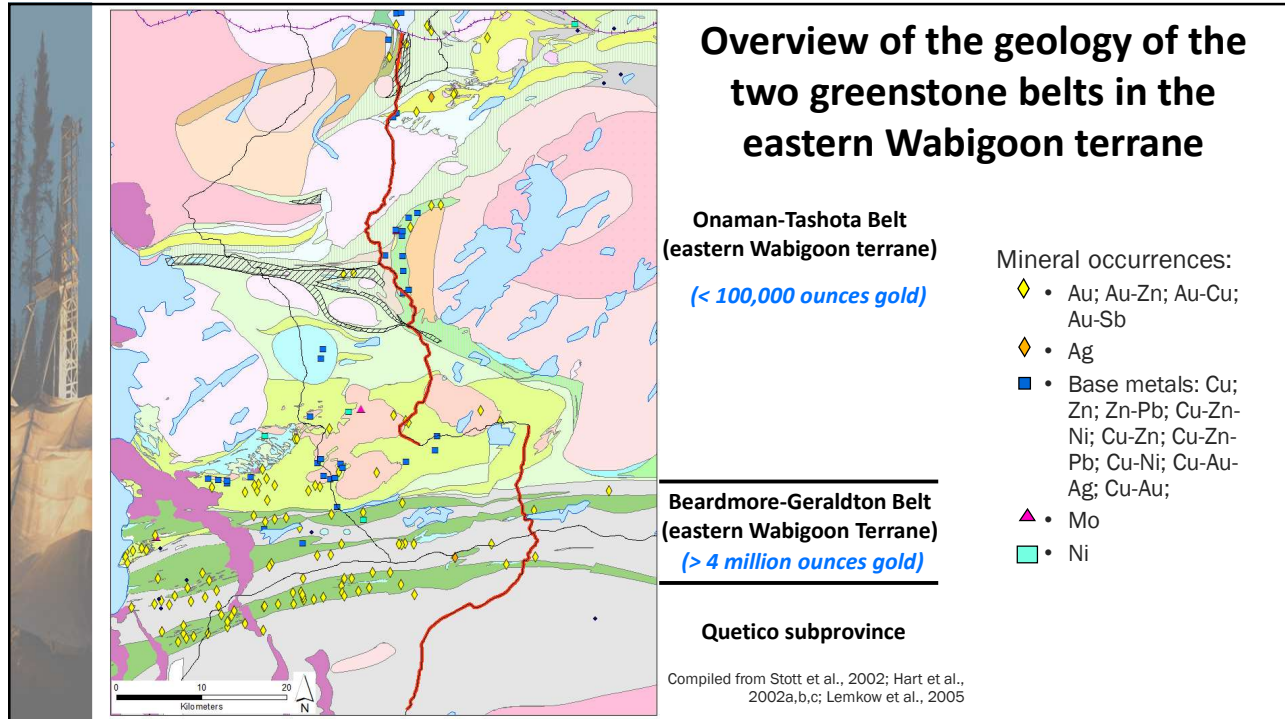
Wabigoon Transects

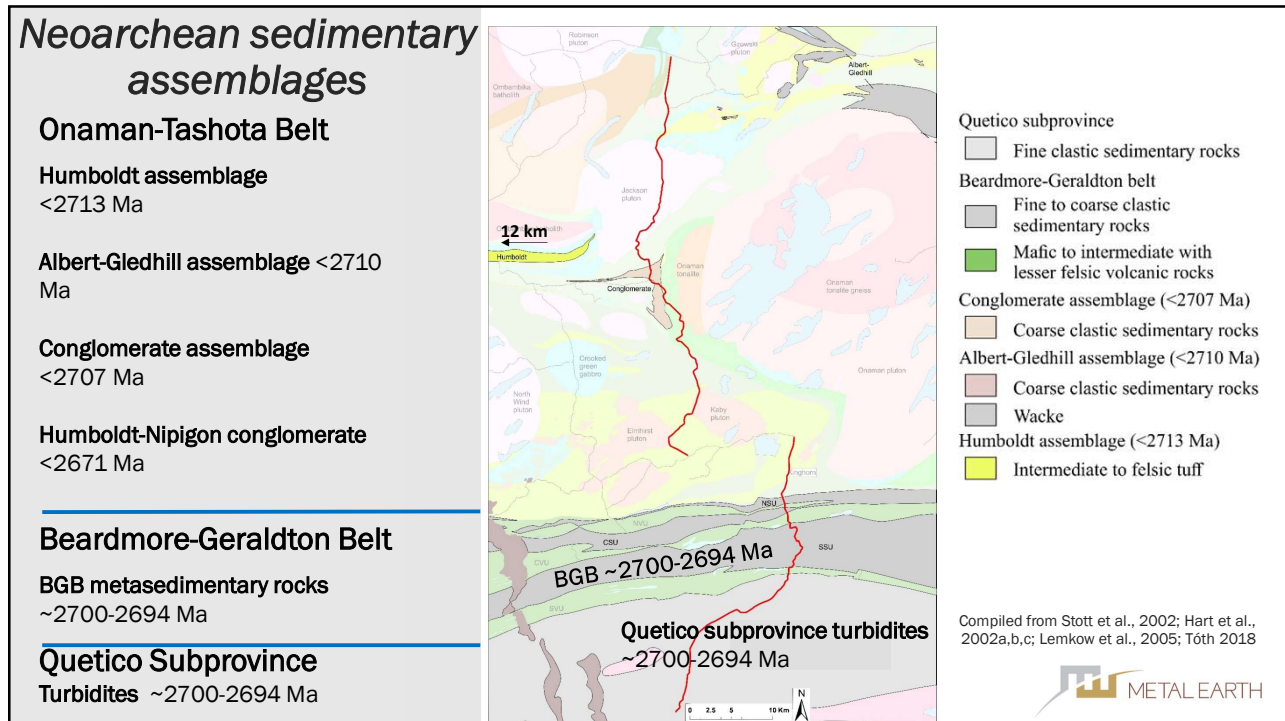
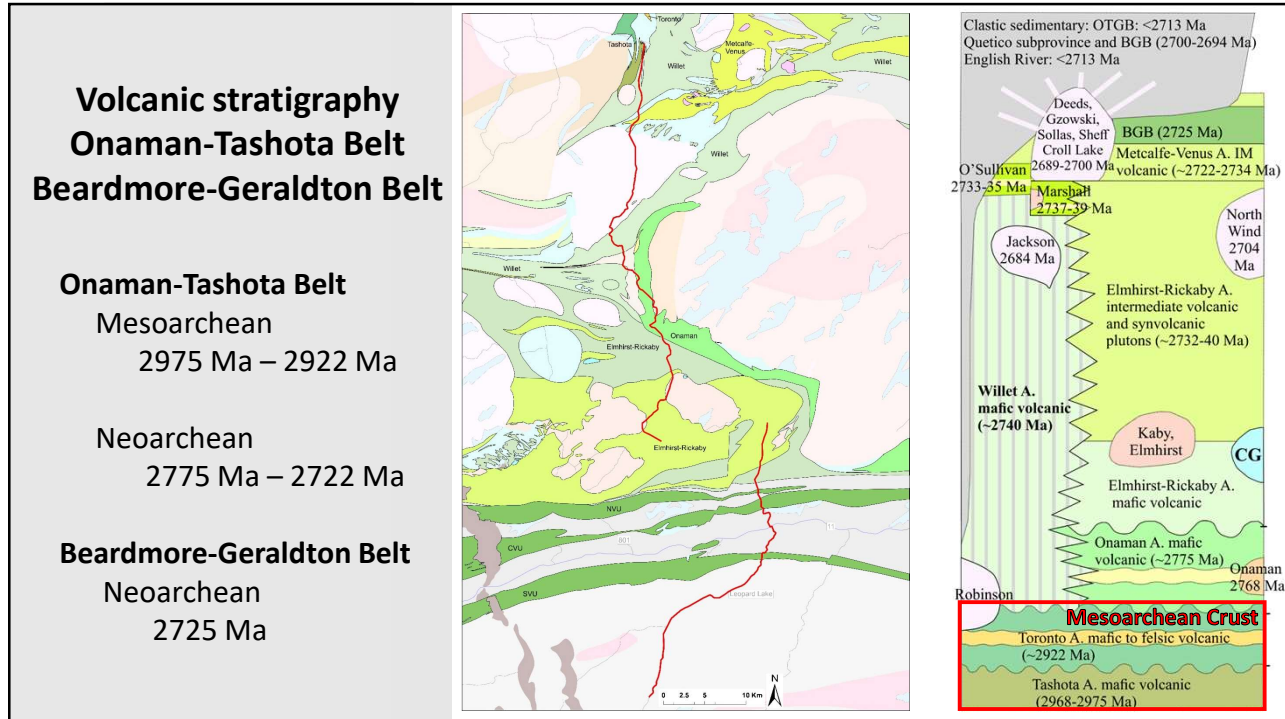


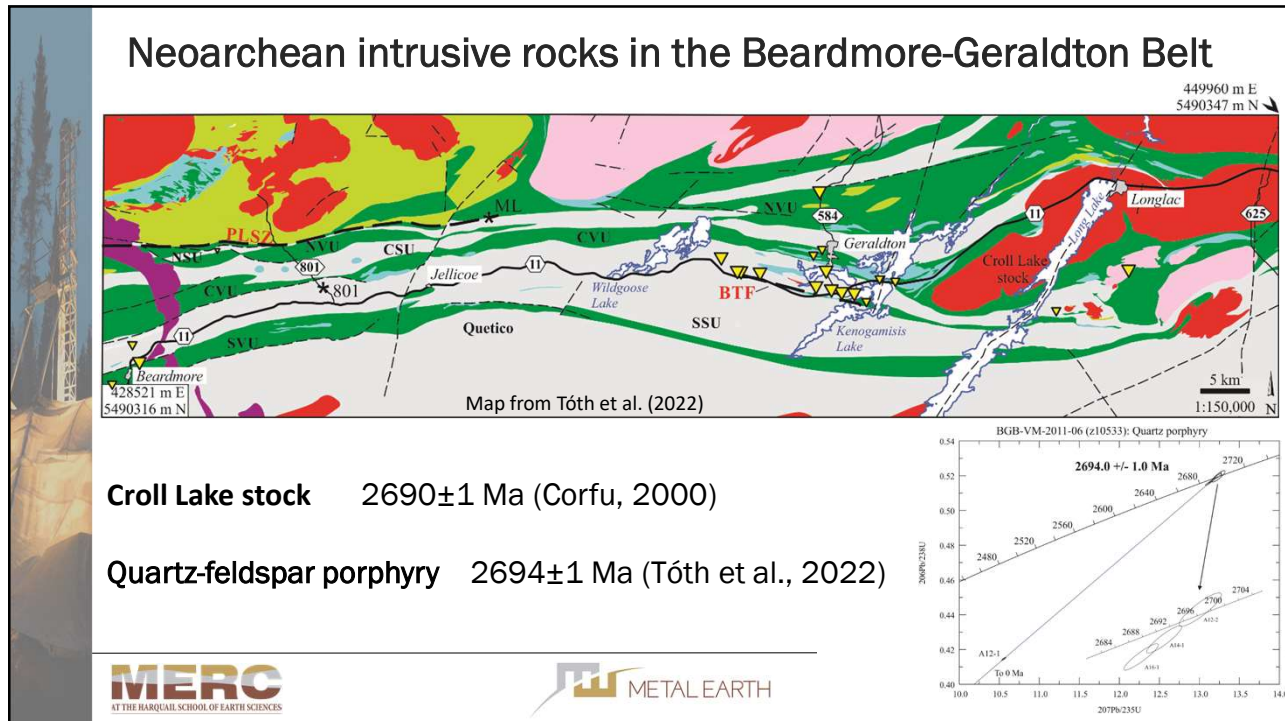
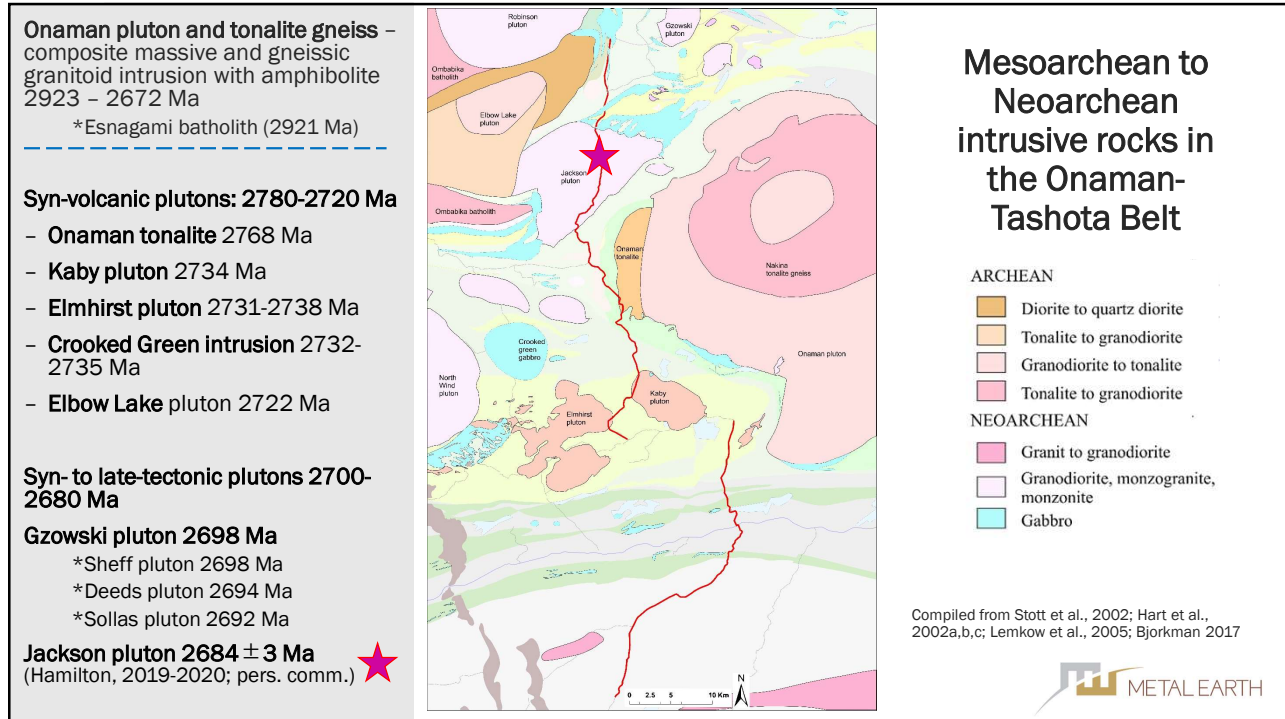
All of this to address the **MAIN QUESTION** of Metal Earth: Why are greenstone belts enriched in metals and others not. What are the processes or factors responsible for the preferential metal endowment of greenstone belts?

OUTLINE

- Overview of the geology of the two main greenstone belts in the eastern Wabigoon subprovince: Onaman-Tashota Belt and Beardmore-Geraldton Belt;
- Comparison of their structural history, including the relative and absolute timing of structures in the two belts;
- Comparison of the gold mineralization history of the two belts;
- Integration of these results with the new seismic and MT transect;
- Summarize the factors and processes responsible for the preferential gold endowment of the eastern Wabigoon subprovince.







Overview of geology of the Onaman-Tashota Belt and Beardmore-Geraldton Belt

	<i>Onaman-Tashota belt</i>	<i>Beardmore-Geraldton belt</i>
Volcanism	Neoproterozoic volcanism (2722- 2780 Ma) Mesoarchean volcanism (2922 Ma -2975 Ma)	Neoproterozoic volcanism (ca. 2725 Ma)
Sedimentation	Turbidites and polymictic conglomerates (2713 Ma - 2692 Ma)	Turbidites and polymictic conglomerates (2700 Ma - 2694 Ma)
Plutonism	Neoproterozoic plutonism Syn-volcanic pluton (2780 Ma – 2720 Ma) Late syn-tectonic plutons (2700 Ma – 2680 Ma) Mesoarchean plutonism (2922 Ma)	Neoproterozoic plutonism Late syn-tectonic plutons (2694- 2690 Ma)

Structural Geology of the Onaman-Tashota Belt and the Beardmore-Geraldton Belt

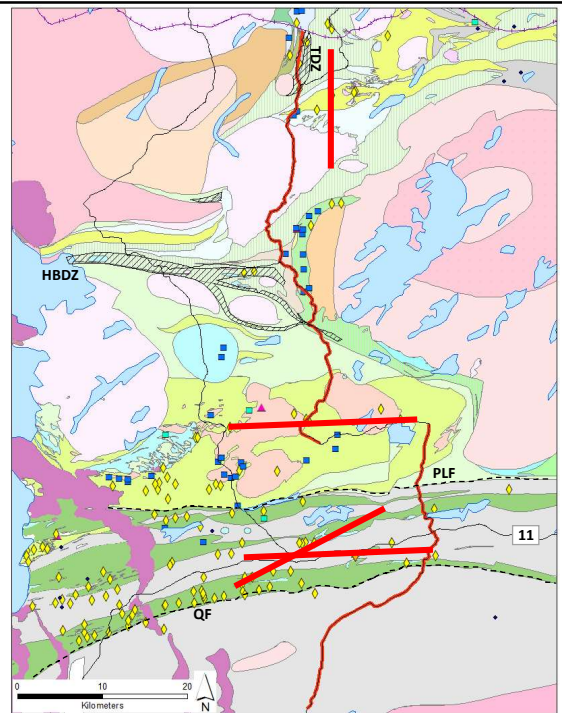
What are the important structures?

Regional foliations

Prominent N-S striking S1 foliation in the central part of the Onaman-Tashota Belt

Prominent E-W striking S2 foliation in the northern and southern part of the Onaman-Tashota Belt, and Beardmore-Geraldton Belt

Prominent NE-SW striking S3 foliation in the Beardmore-Geraldton Belt



Compiled from Stott et al., 2002; Hart et al., 2002a,b,c; Lemkow et al., 2005; OGS MDI July, 2018

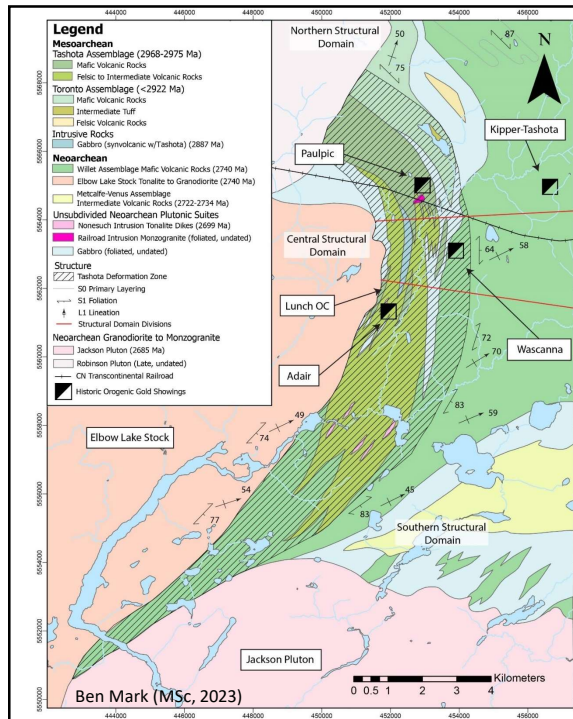
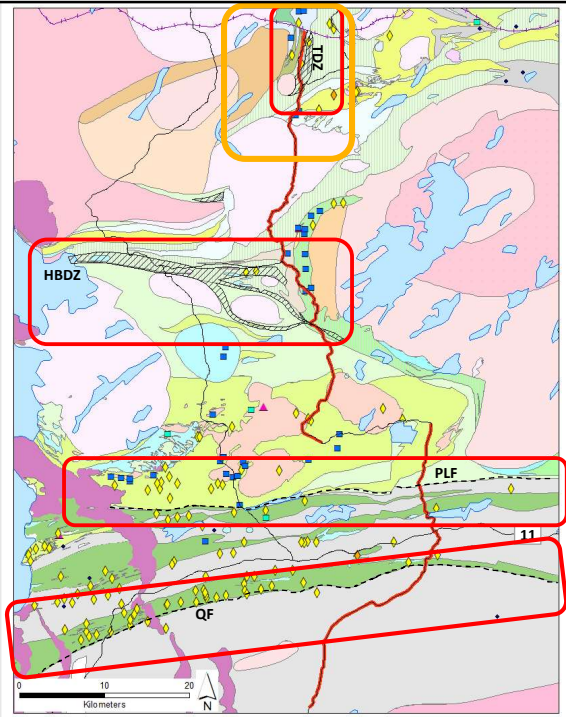
Structural Geology of the Onaman-Tashota Belt and the Beardmore-Geraldton Belt

What are the important structures?

Deformation Zones

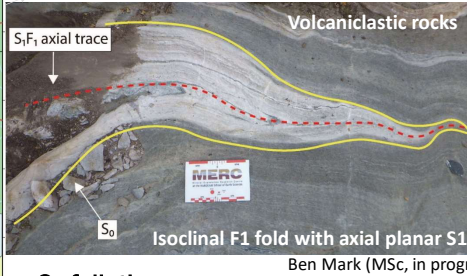
- D1 Tashota deformation zone (TDZ)**
 - N-striking deformation along batholith-volcanic contact
- D2 & D3 Humboldt Bay deformation zone (HBZD)**
 - E-striking deformation corridor within the OTB
- D2 & D3 Paint Lake Fault (PLF)**
 - E-striking deformation zone separating the BGB from the OTB
- D2 & D3 Quetico Fault (QF)**
 - E-striking deformation zone separating the BGB from the Quetico

Compiled from Stott et al., 2002; Hart et al., 2002a,b,c; Lemkow et al., 2005; OGS MDI July, 2018



Structural Geology of the Onaman-Tashota Belt

D1 deformation event Tashota deformation zone



S₁ foliation

- Dominant foliation in the central part of the OTB, where it wraps around older Mesoarchean and Neoarchean syn-volcanic intrusions

L₁ lineation

- Steep mineral and stretching lineation along S₁ and defined by mineral streaks or elongate clasts

D₁ deformation event – Age constraints

Quartz-feldspar porphyry and tonalite dikes contain S₁ foliation in the Humboldt Bay deformation zone and Tashota deformation zone, respectively

Crystallization age of dikes:

HBDZ: 2699.1 ± 1.7 Ma TDZ: 2699.5 ± 1.6 Ma

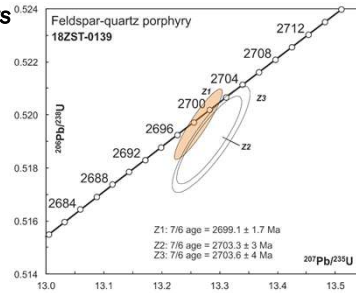
Maximum age of D₁ is 2699 Ma.

Jackson pluton granodiorite crosscuts S₁ foliation and is itself not foliated

Crystallization age of Jackson pluton: 2684 ± 3 Ma

Minimum age of D₁ is 2684 ± 3 Ma.

D₁ occurred between 2699 Ma and 2684 Ma

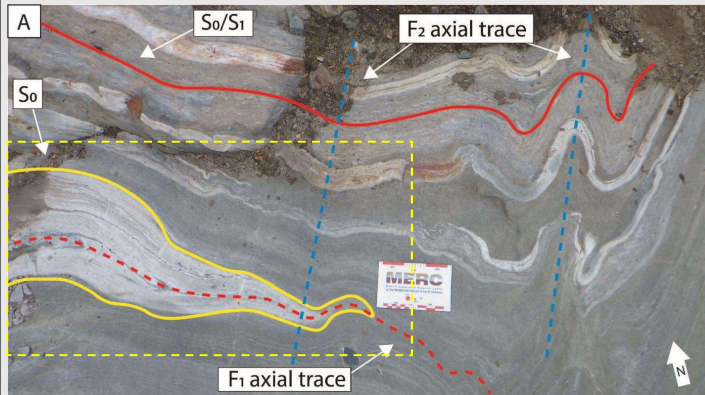


U-Pb ages: Hamilton, 2019-2020; pers. comm.



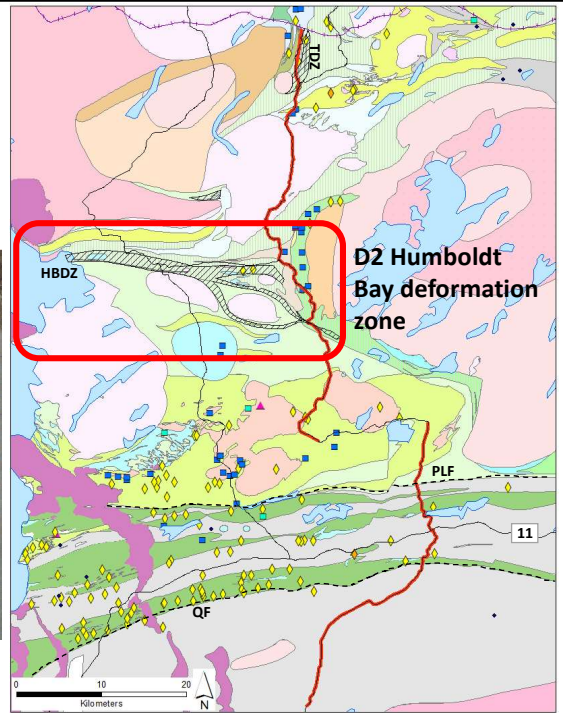
Structural Geology of the Onaman-Tashota Belt D₂ deformation event

Lunch outcrop in D₁ Tashota deformation zone



Ben Mark (MSc, 2023)

Compiled from Stott et al., 2002; Hart et al., 2002a,b,c; Lemkow et al., 2005; OGS MDI July, 2018

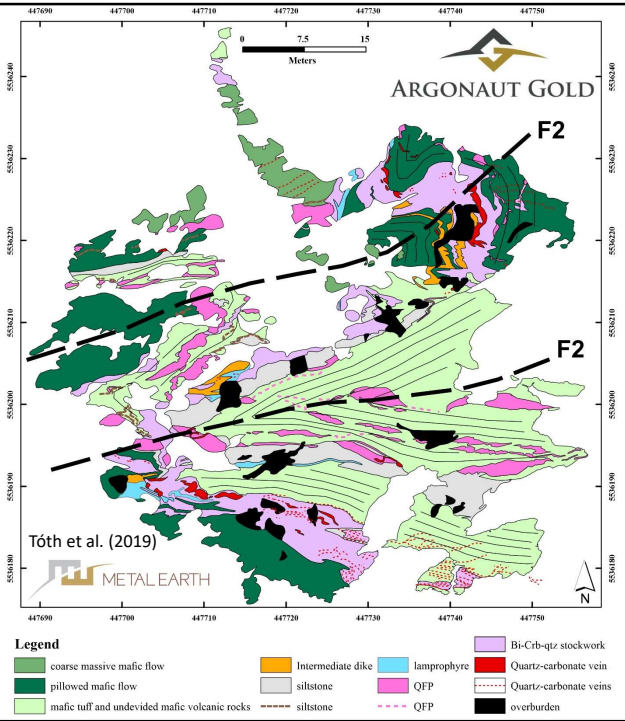
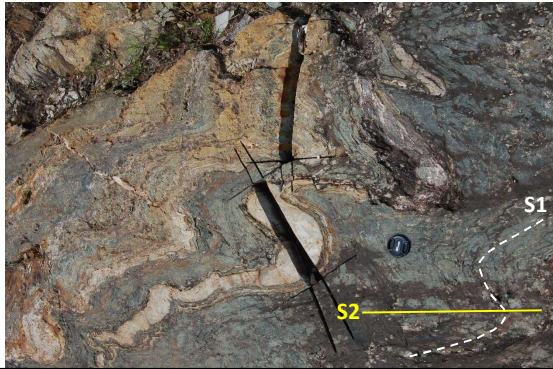


D2 deformation event in OTB

N-S striking S1 foliation is transposed parallel to EW-striking F2 folds and S2 foliation along the Humboldt Bay deformation zone.

S2 foliation and F2 folds become the most prominent structures from the Humboldt Bay deformation zone to the Paint Lake Fault along the southern boundary of the Onaman-Tashota Belt.

D2 bracketed between 2699 Ma and 2667 Ma ($^{40}\text{Ar}-^{39}\text{Ar}$ ages from amphibolite; Culshaw et al. 2006).



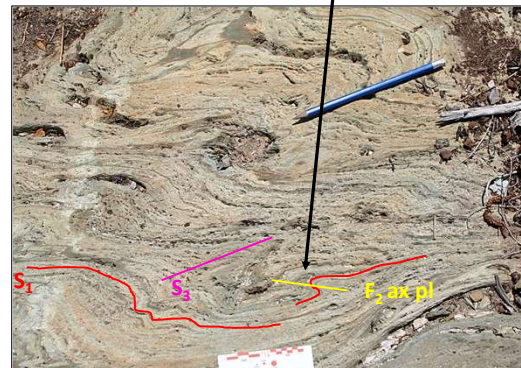
D3 deformation event- OTB

Expressed by Z-shaped F3 folds with an axial plane slaty cleavage or crenulation cleavage

EW-striking dextral faults or shear zones

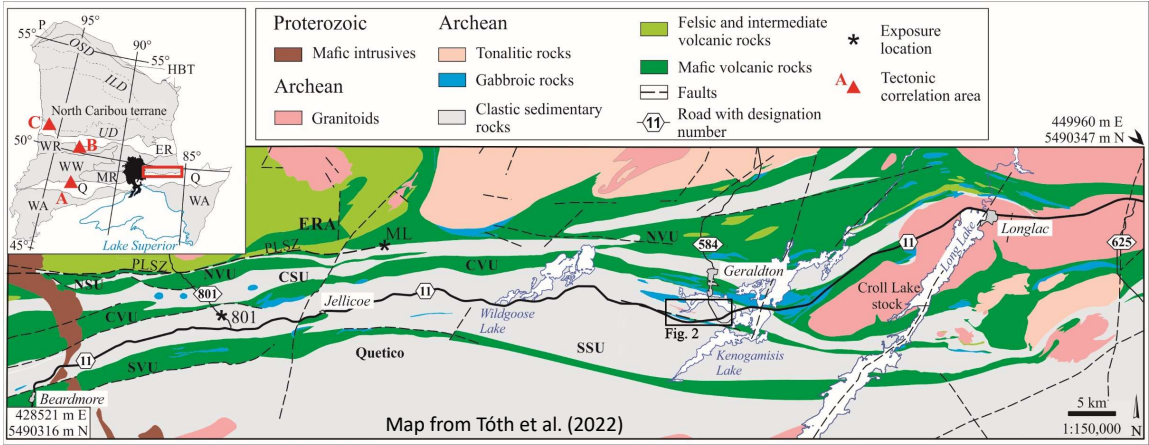
< 2667 Ma in age

Ben Mark (MSc, 2023)

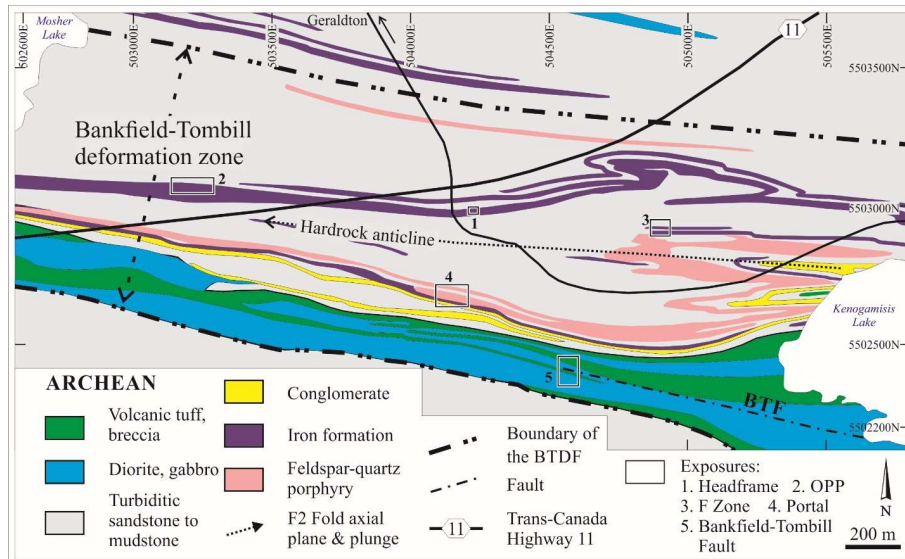


Structural Geology of the Beardmore-Geraldton Belt

The BGB consists of interleaved panels of sedimentary rocks and volcanic rocks, which are cut by the Croll Lake stock at the eastern end of the belt.



All D1 to D3 structures are observed in the Geraldton area within the 1 km wide Bankfield-Tombill deformation zone.

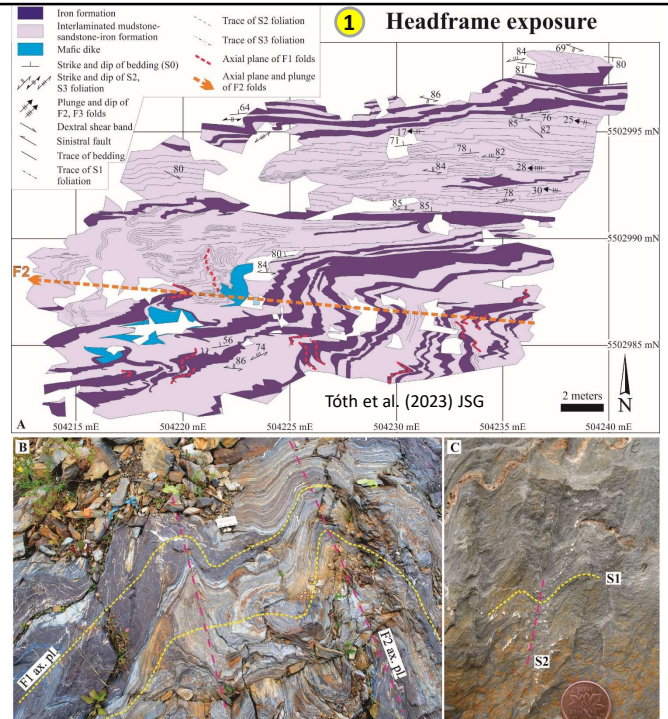
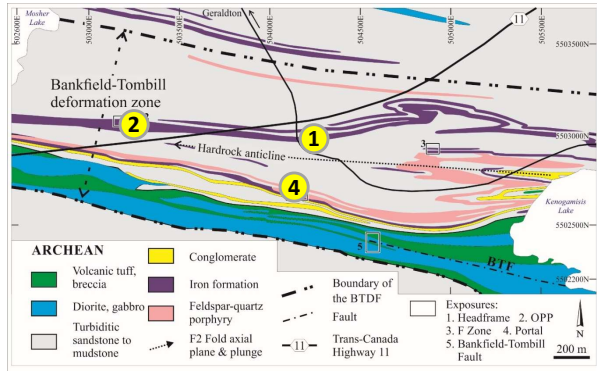


D1 deformation event

Isoclinal F1 folds with axial planar S1 foliation

Bedding-parallel in iron formation and sedimentary rocks

Observed in mafic dikes and QFP intrusions

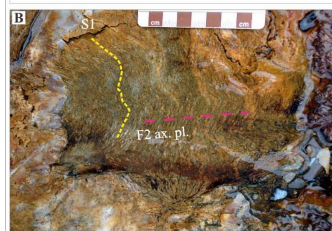
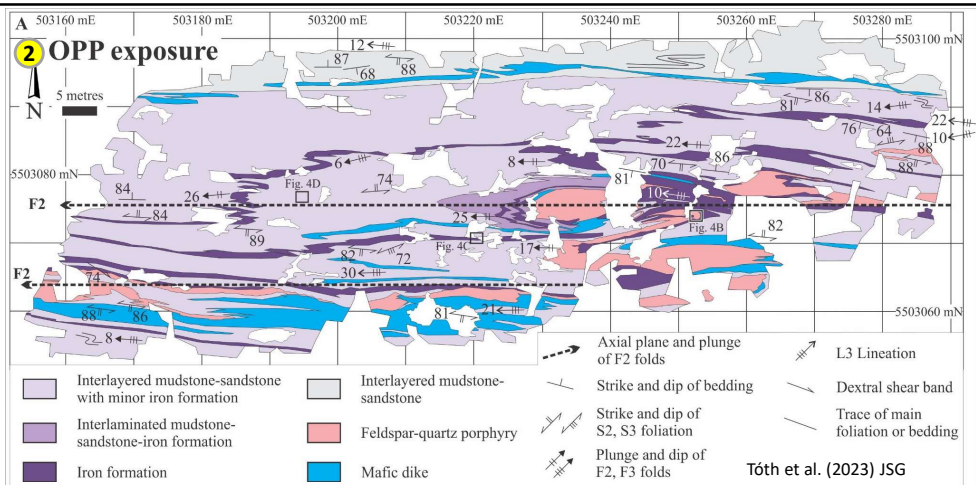


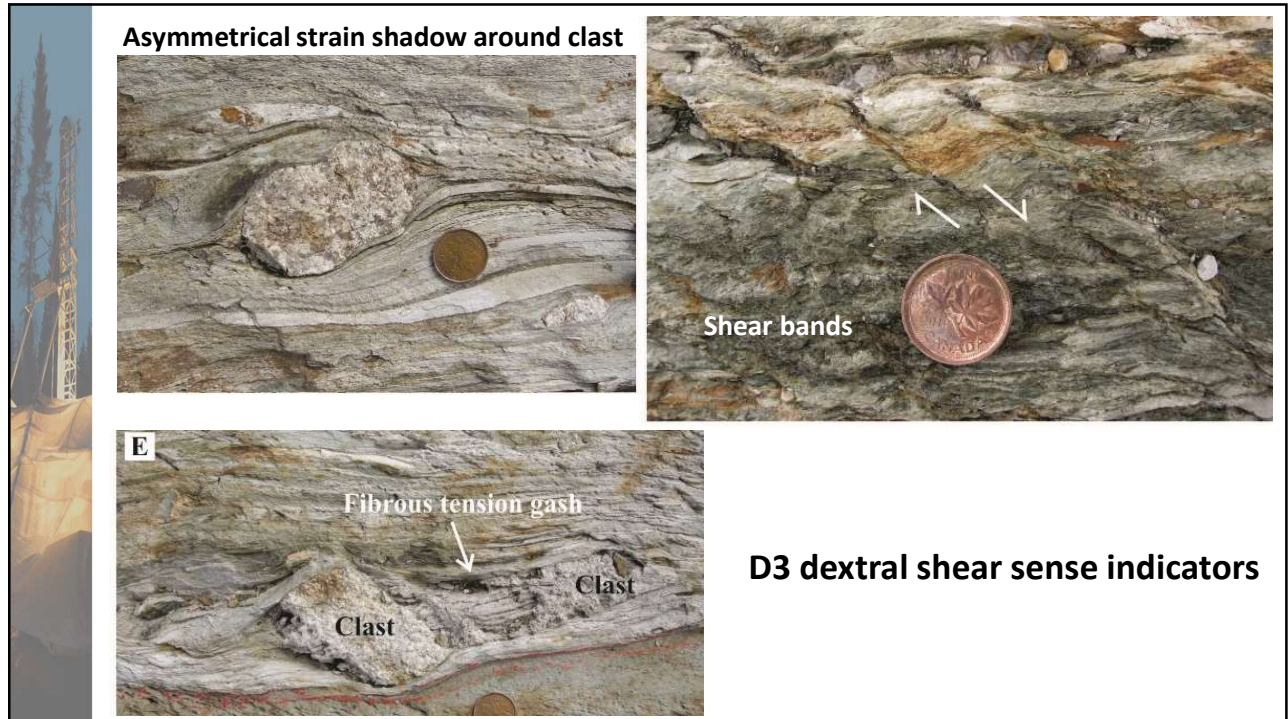
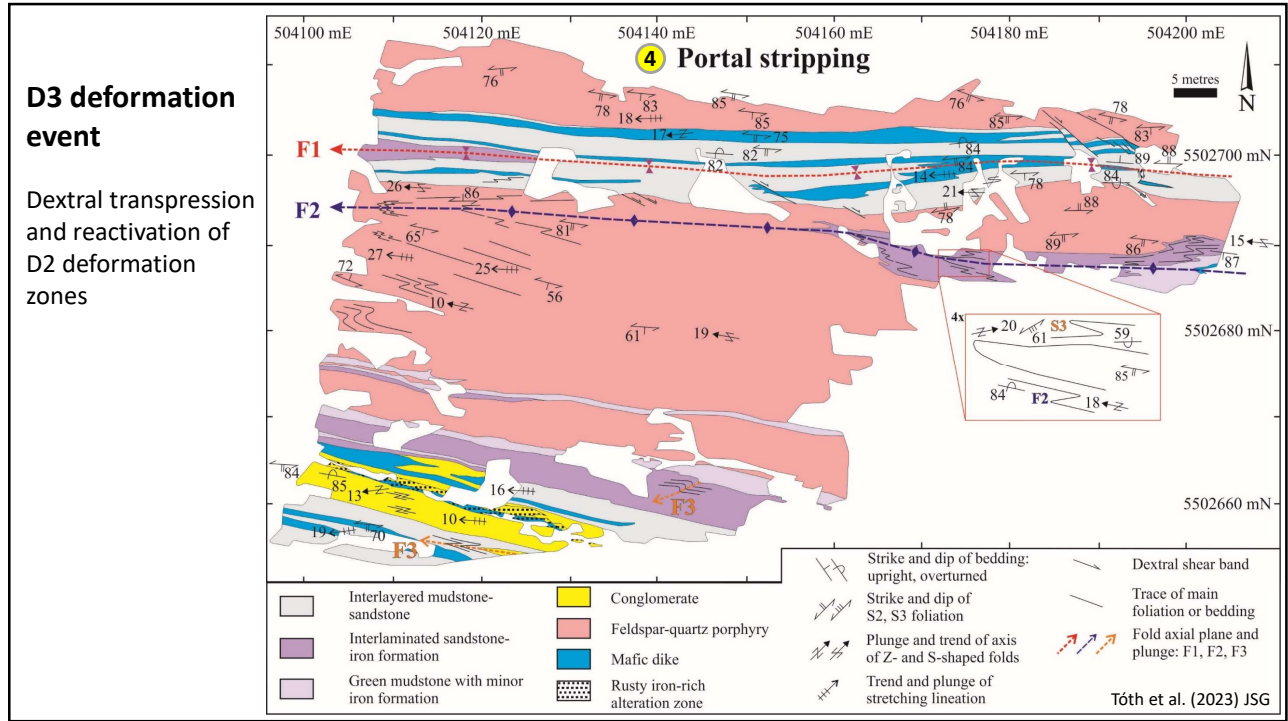
D2 deformation event

Development of D2 deformation zone and regional S2 foliation

Refolding and transposition of F1 folds and S1 foliation

Observed in mafic dikes and QFP intrusions

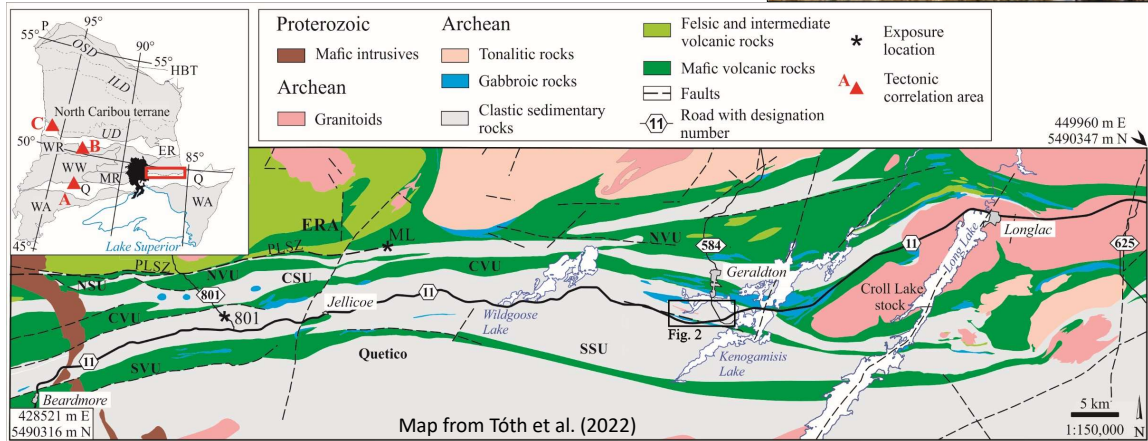
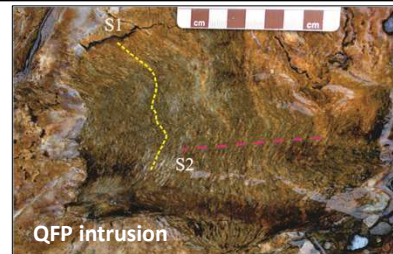




Chronological constraints on the deformation events in the Beardmore-Geraldton Belt

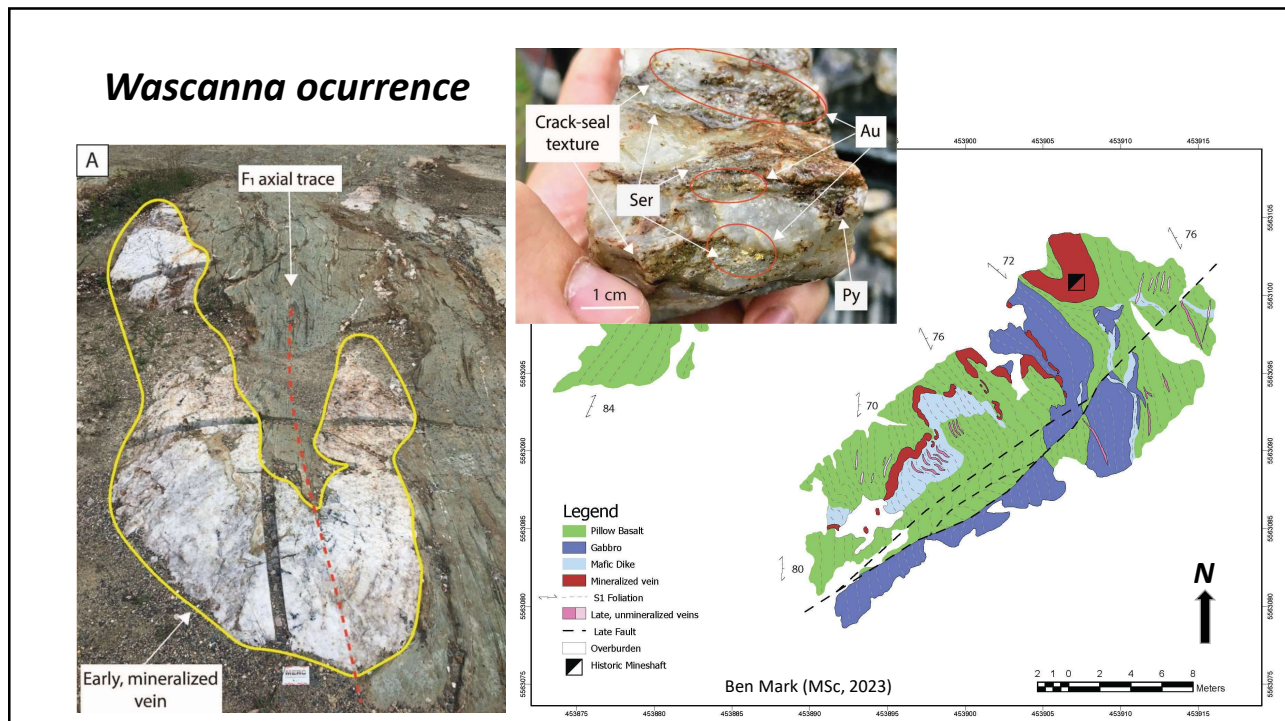
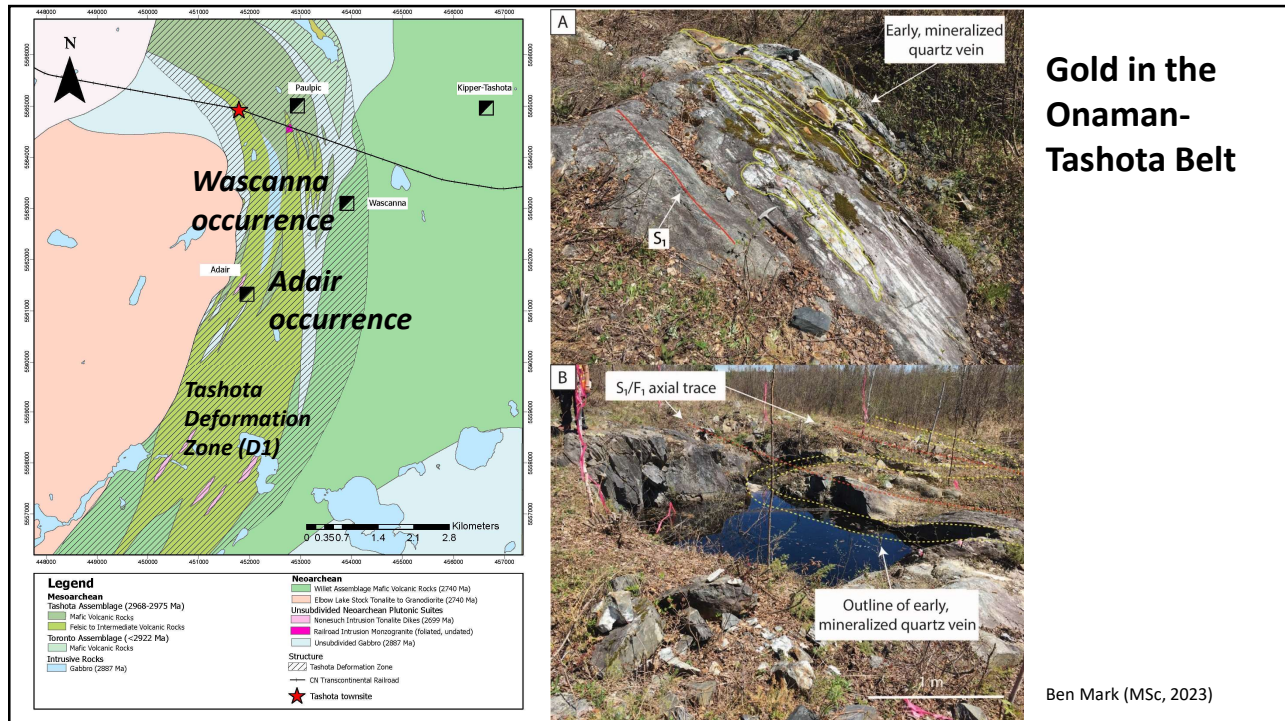
D1 is bracketed between 2694 Ma, S1-foliated QFP intrusion, and 2690 Ma, the age of the Croll Lake stock.

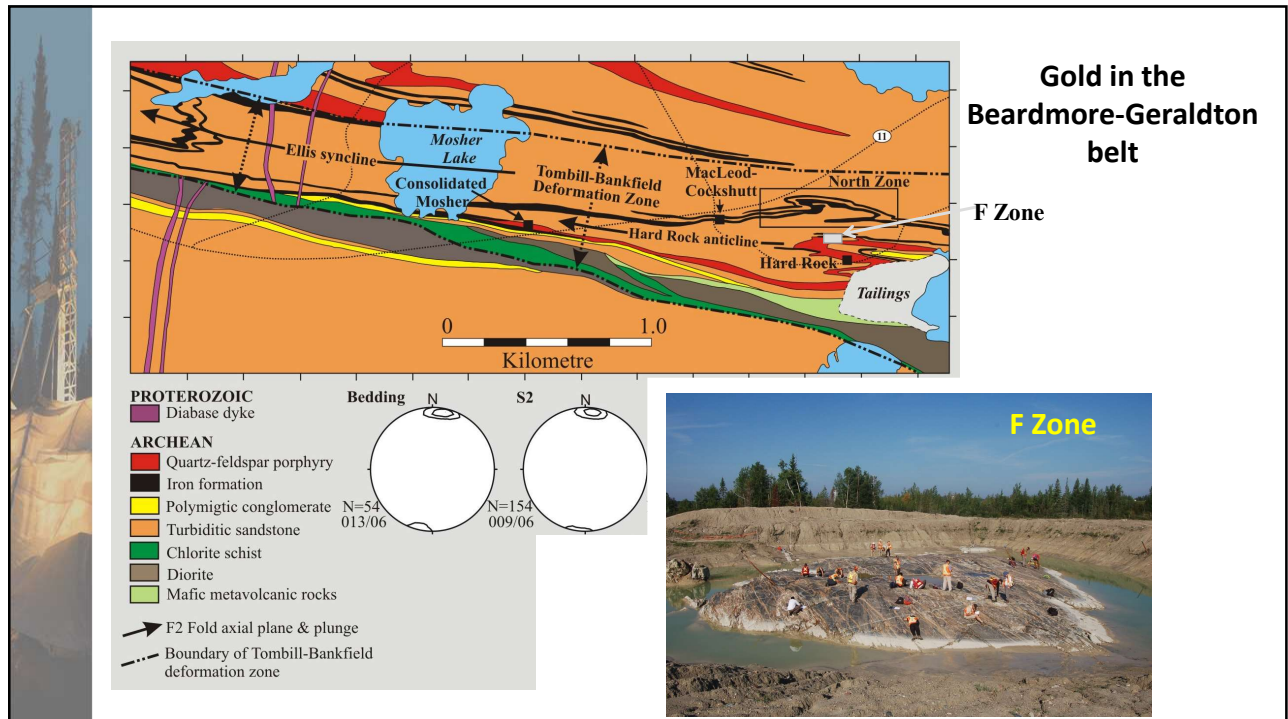
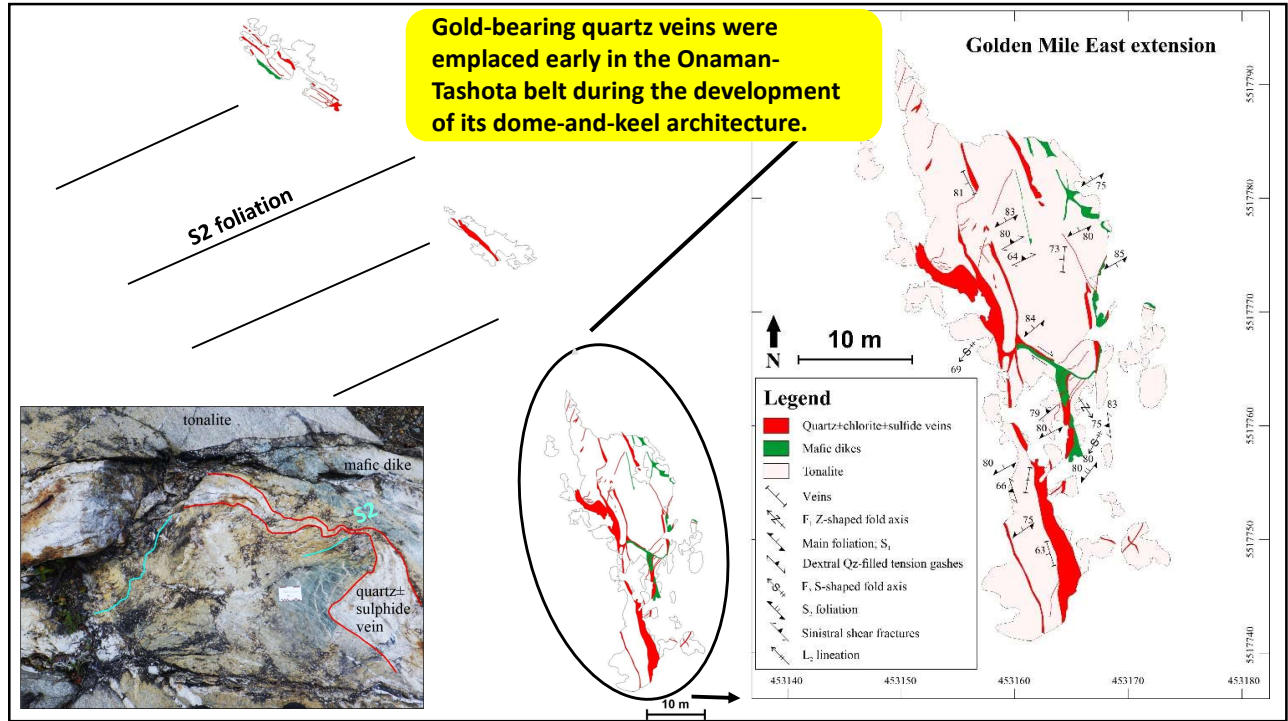
D2 and D3 events must be less than 2690 Ma.

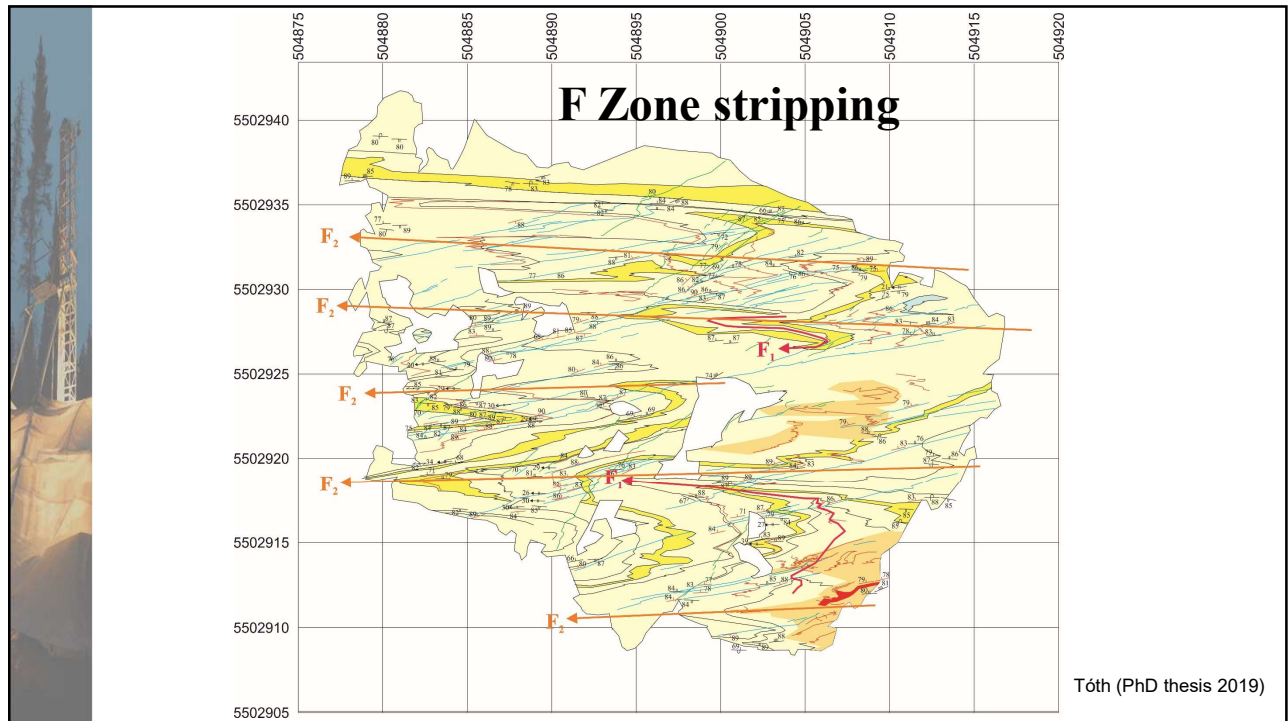


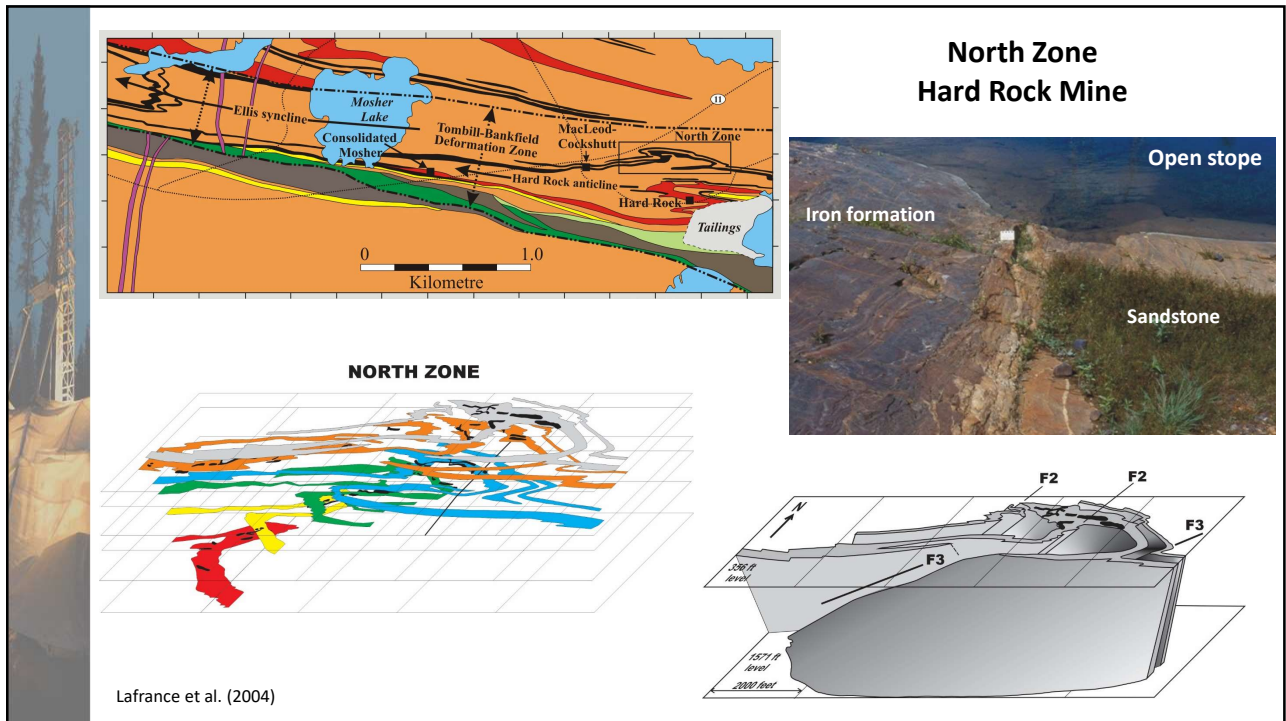
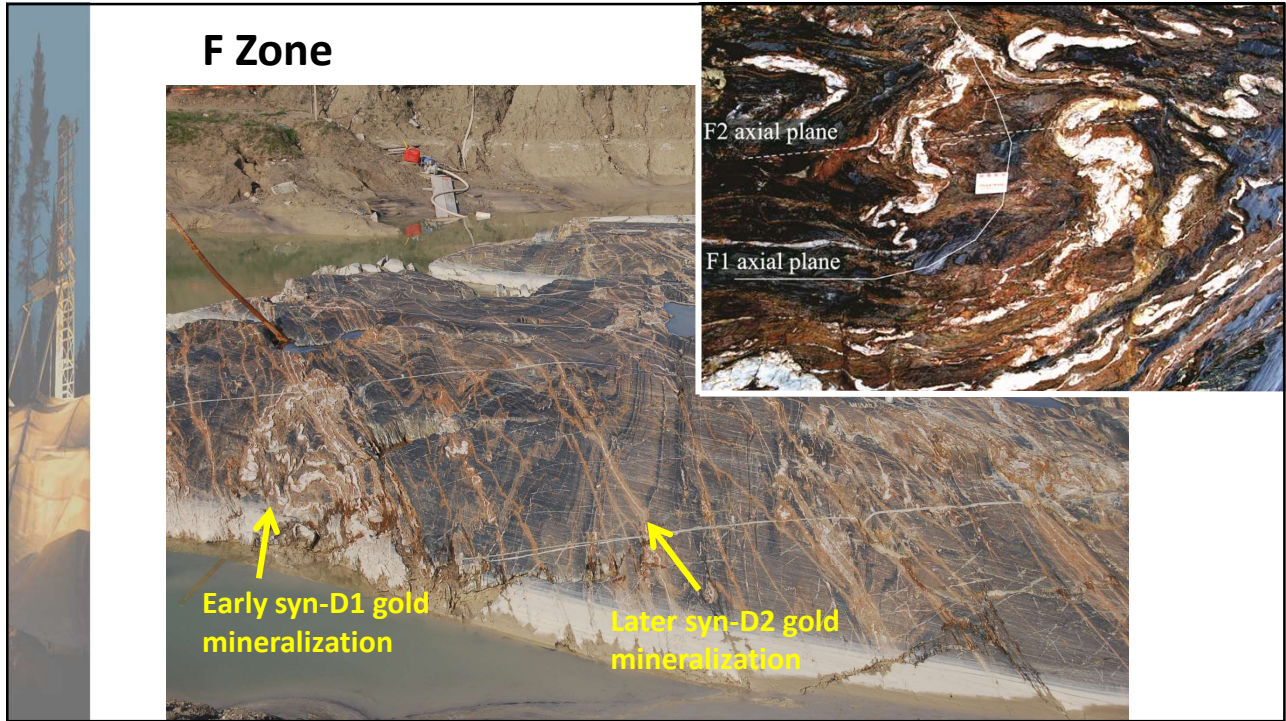
Comparison of structural geology of the OTB and the BGB

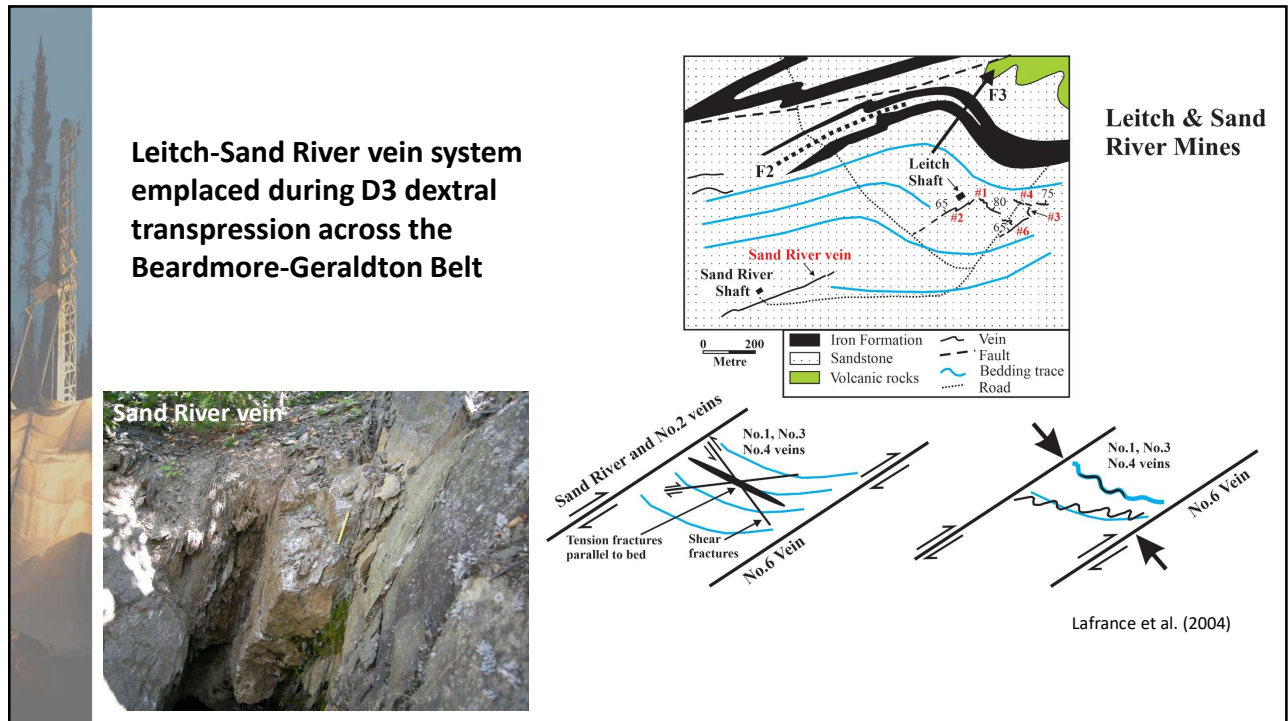
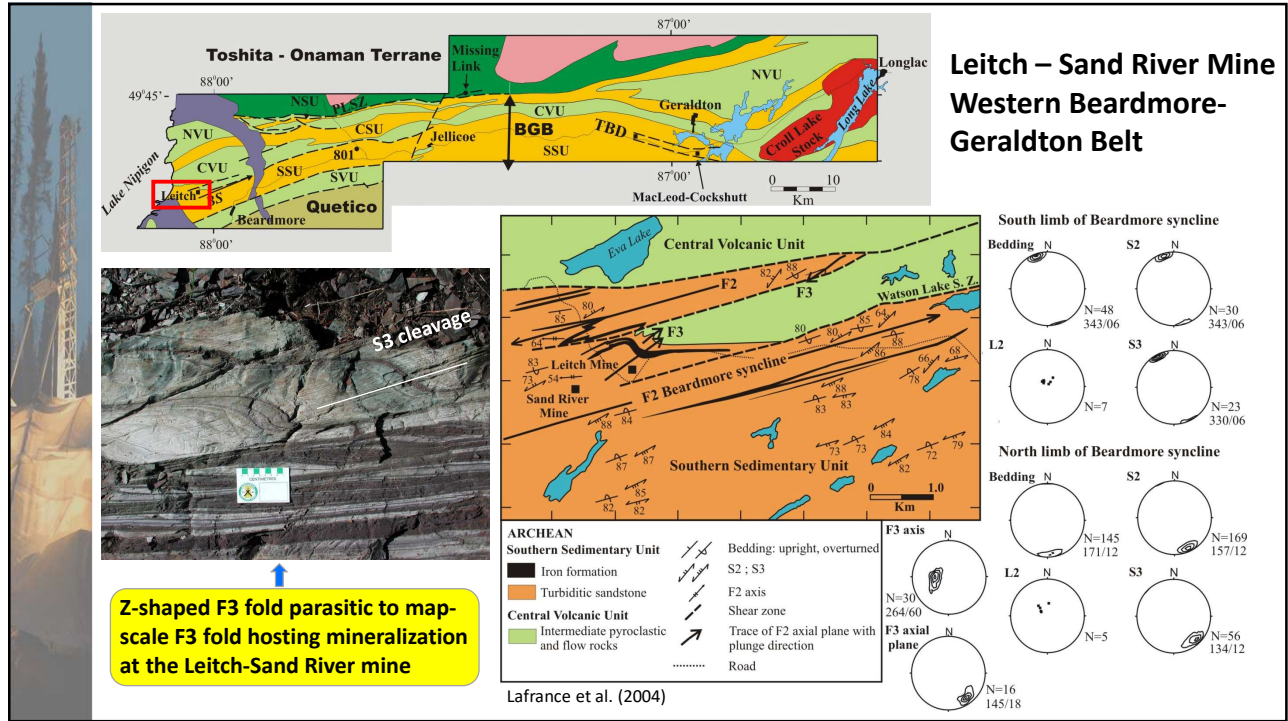
	<i>Onaman-Tashota belt</i>	<i>Beardmore-Geraldton belt</i>
D₁	Formation of dome-and-keel architecture characterized by the formation of granitoid intrusions and folding of the dome	Formation of linear accretionary belt by thrust character of the belt, formation of sedimentary panels
	Coeval development of the dome-and-keel Onaman-Tashota belt and linear accretionary Beardmore-Geraldton belt	
	<2699.1 ± 1.7 Ma - 2684 Ma	2694+/-1 Ma - 2690 Ma
D₂	Regional fold and E-striking S ₂ foliation D2 deformation zones (e.g. Tombill-Bankfield deformation zone)	Regional fold and E-striking S ₂ foliation D2 deformation zones (e.g. Tombill-Bankfield deformation zone)
	Same	
	2699 Ma - 2667 Ma	<2690 Ma
D₃	Z-shaped F3 folds and minor E-striking dextral transcurrent faults	Dextral shearing, reactivation of D2 deformation
	More significant in BGB	
	<2667 Ma	











Comparison of gold occurrences in Onaman-Tashota Belt and Beardmore-Geraldton Belt

	<i>Onaman-Tashota belt</i>	<i>Beardmore-Geraldton belt</i>
<i>Structural association</i>	Deformation zones	Deformation zones and regional fold hinges
<i>Structural chronology</i>	D1 dome-and-keel development	D1 accretion along the southern margin of the Wabigoon subprovince D2 deformation zones and fold hinges D3 dextral transpression
<i>Total production</i>	<100,000 ounces gold	> 4 million ounces gold



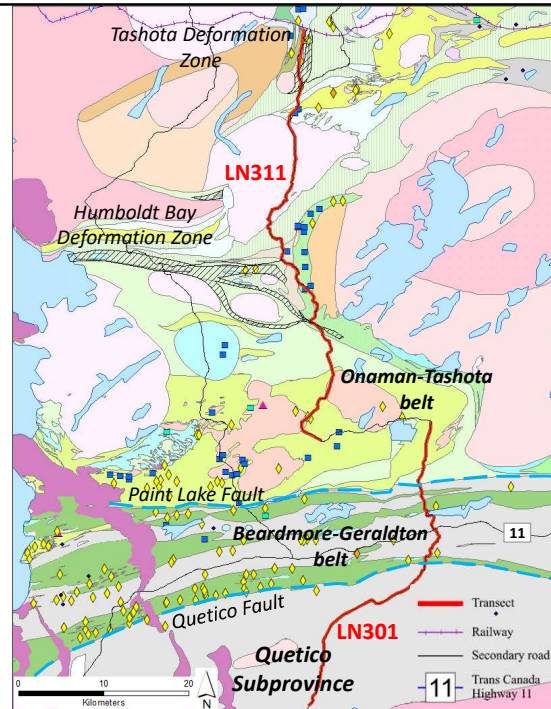
Seismic and Magnetotelluric (MT) Transect

Transect Lines in Red

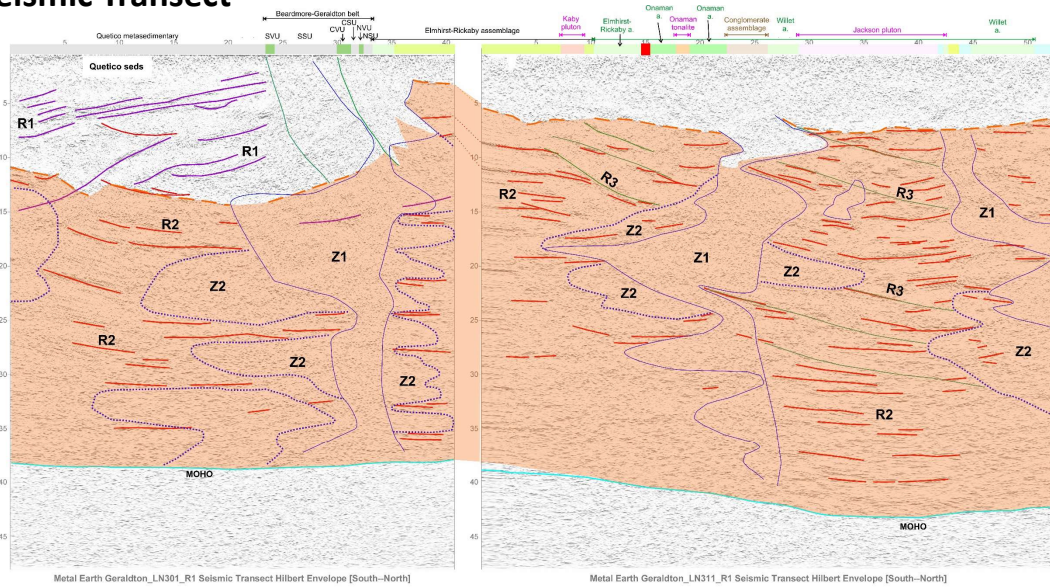
LN301 and the first 25 km of LN311 cut across east-trending stratigraphy.

The last 25 km of LN311 are oblique and at 45° to stratigraphy.

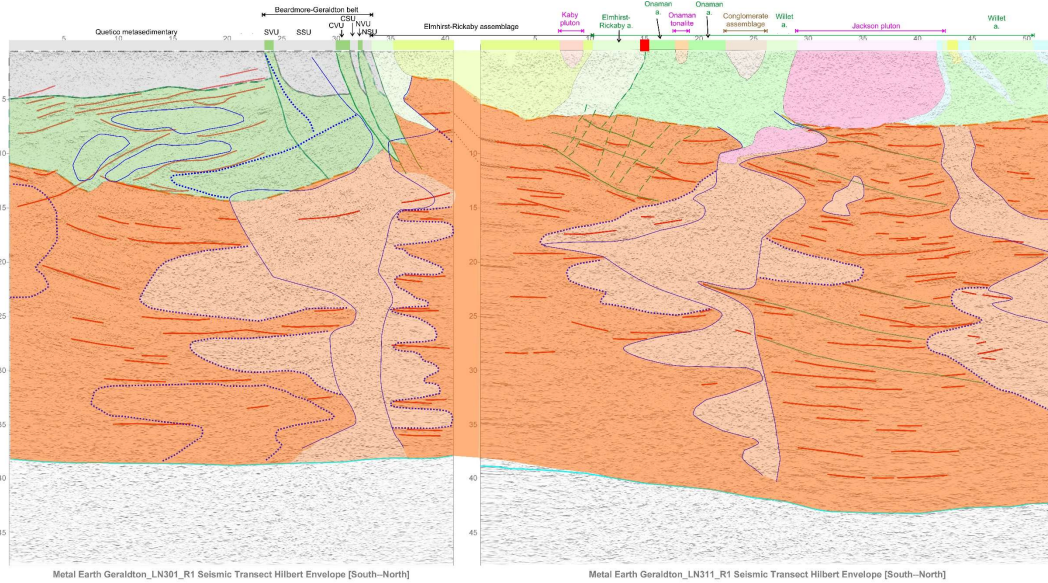
Compiled from Stott et al., 2002; Hart et al., 2002a,b,c; Lemkow et al., 2005; Bjorkman 2017

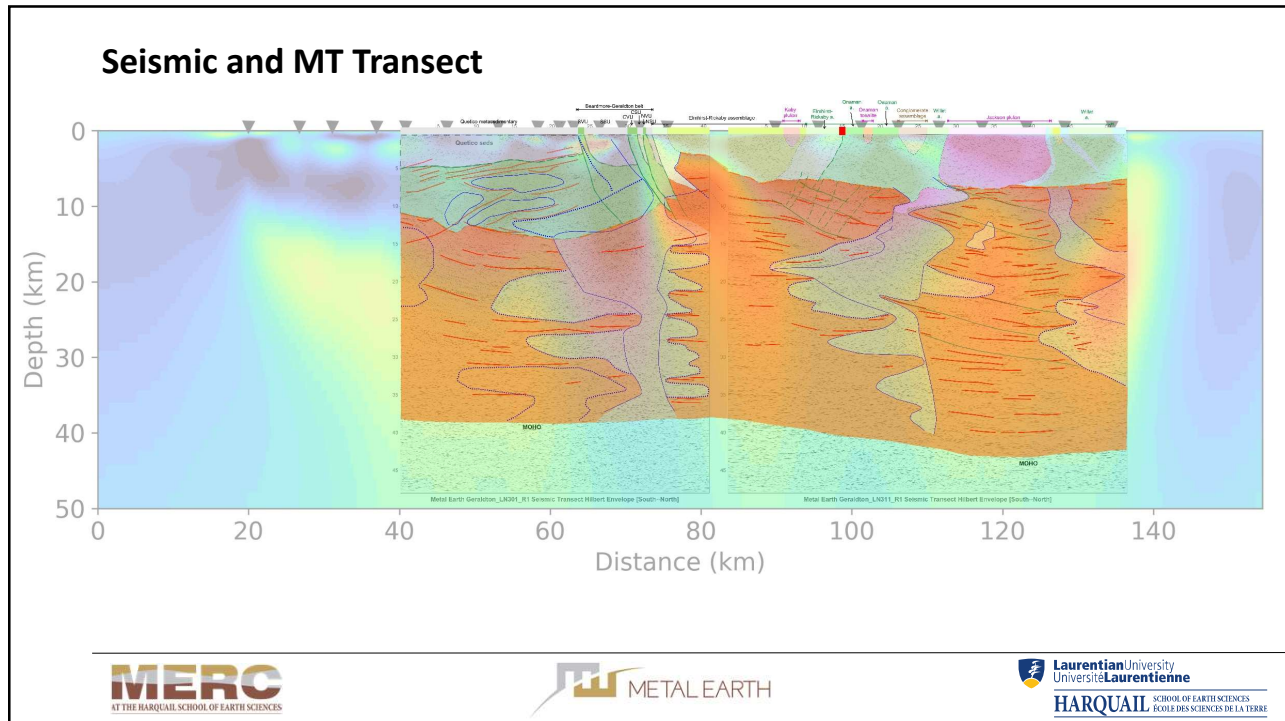
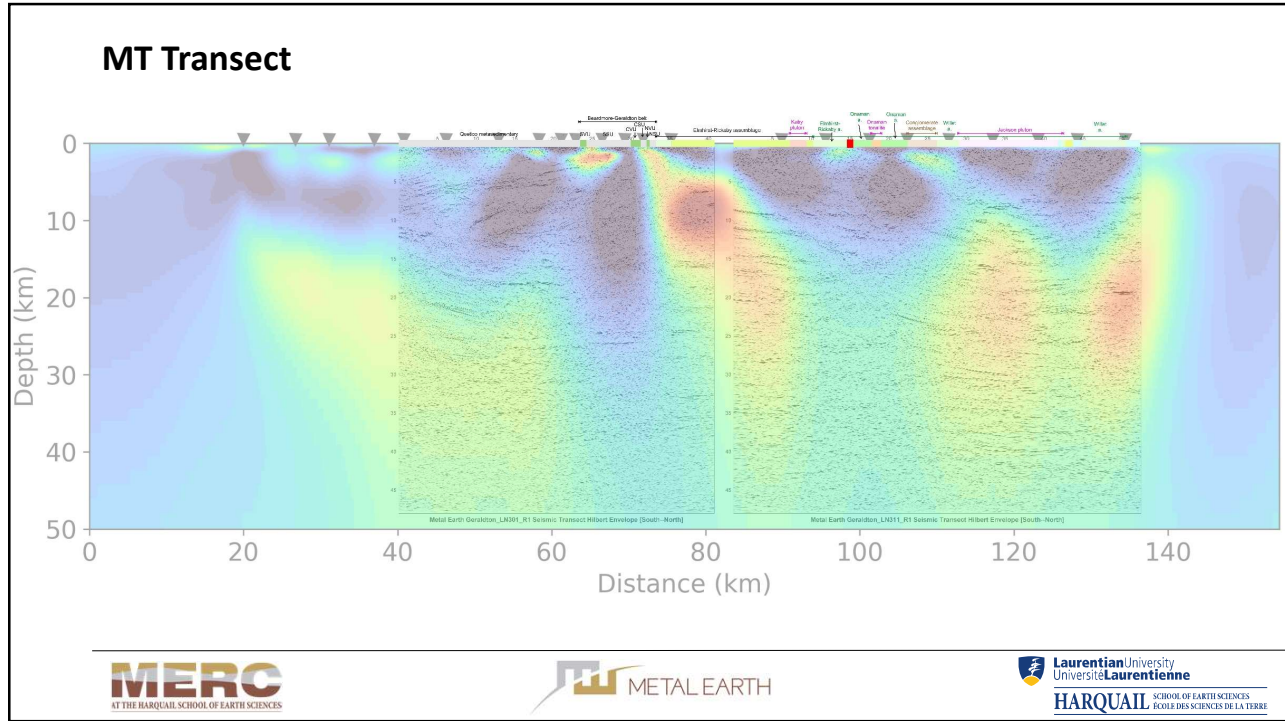


Seismic Transect



Interpreted Seismic Transect



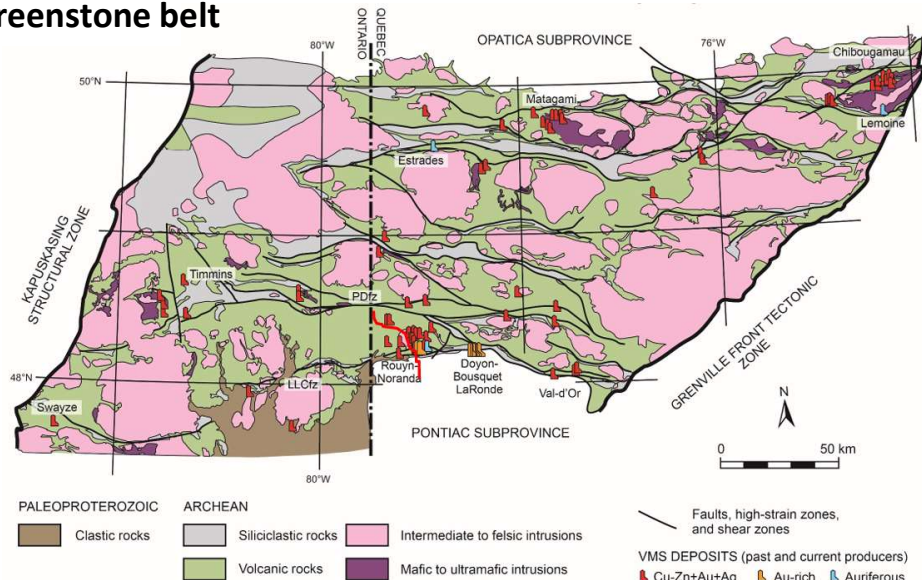


Key Take-Aways

- (1) The development of the **dome-and-keel** architecture in the Onaman-Tashota Belt was **coeval** with the development of the **linear accretionary** architecture of the Beardmore-Geraldton belt.
- (2) Gold mineralization was emplaced early during the development of their dome-and-keel and linear accretionary architecture.
- (3) The more gold-endowed Beardmore-Geraldton Belt coincides with more conductive (or less resistive) and less reflective steeply-dipping zone(s) on the combined Seismic-MT transect.
- (4) The gold-endowed Beardmore-Geraldton Belt differs from the less-endowed Onaman-Tashota Belt by the presence of multiple steeply-dipping penetrating structures and their reactivation during multiple gold mineralizing events.

Linear accretionary belts are more prospective than dome-and-keel belts!

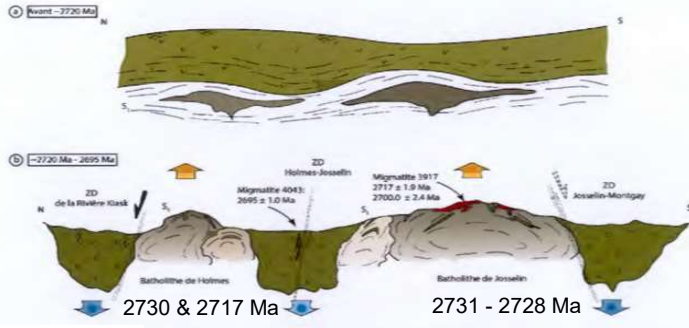
Abitibi greenstone belt



Modified from Monecke et al. (2017) – *Reviews in Economic Geology*, v. 19, p. 7-49

Structural evolution of the northern Abitibi

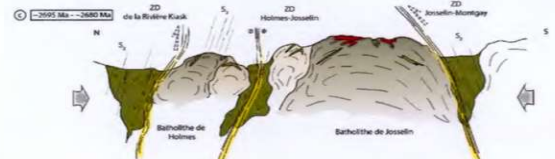
Lebel-sur-Quévillon Attic Complex



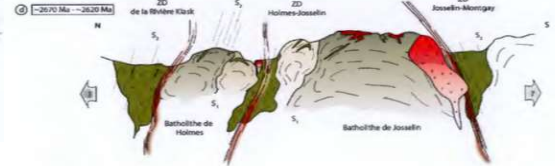
D1-D2 Dome-and-keel formation
 ≤ 2720 to 2695 Ma

Nicolas Revelli PhD thesis 2021 UQAM

D3 Formation of EW regional folds and foliation
 ≤ 2695 to 2680 Ma



D4 Dextral transcurent faulting 2670 to 2620 Ma

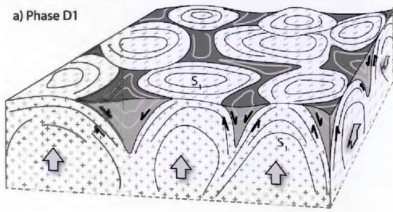


Légende
 Roches volcaniques
 TTG
 Migmatites
 Granites
 Structures en coussins
 S₁
 S₂

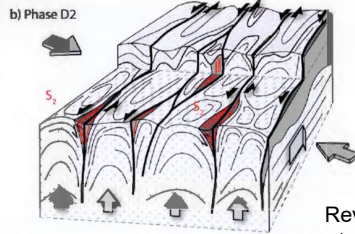


Structural evolution of the northern Abitibi

Lebel-sur-Quévillon Attic Complex



D1-D2 Dome-and-keel
 ≤ 2720 to 2695 Ma

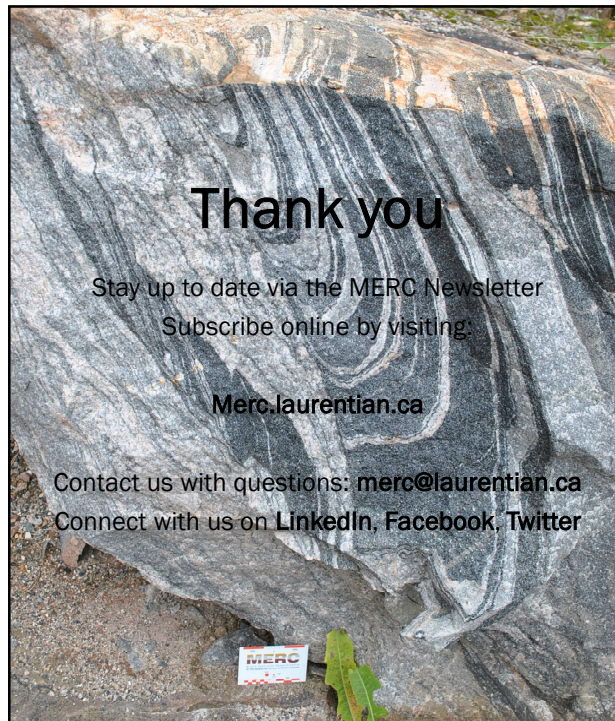
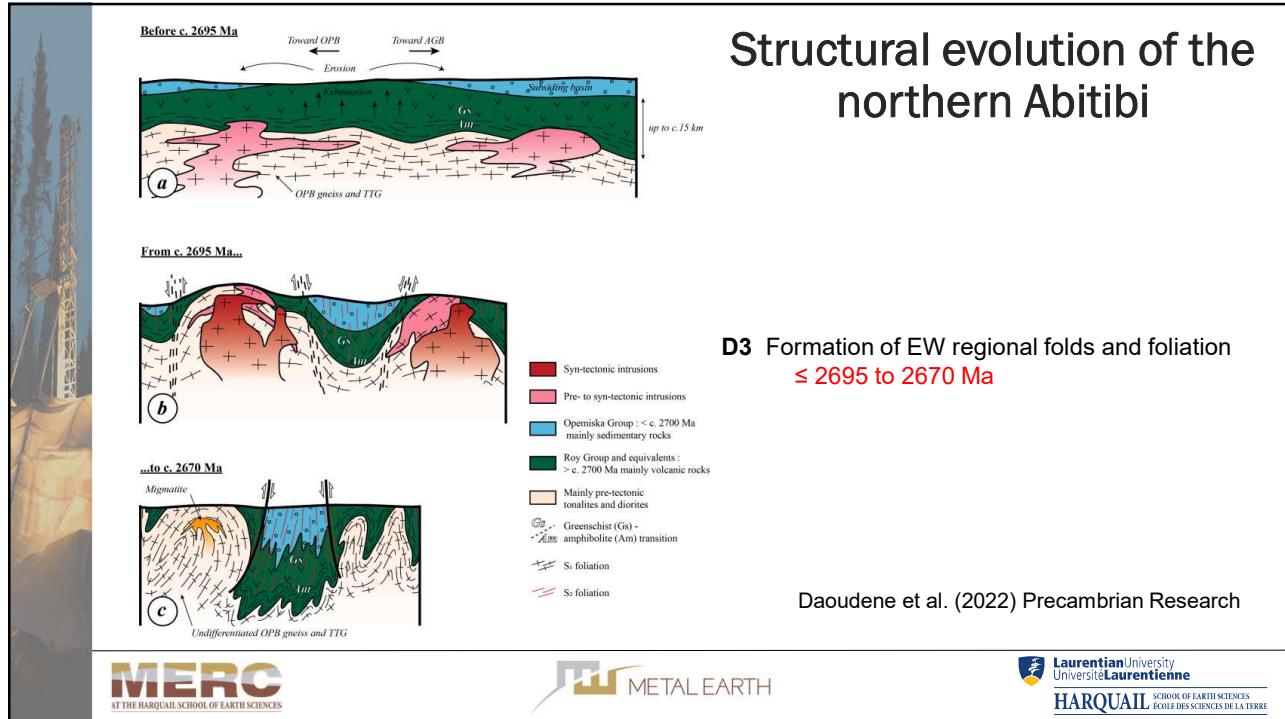


D3 Formation of EW regional folds and foliation
 ≤ 2695 to 2680 Ma

TTG
 Roches volcaniques

Revelli (2021) PhD thesis UQAM *Diagram after Vidal et al. (2009)*





A new Canadian research initiative funded by Canada First Research Excellence Fund.

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 ÉCOLE DES SCIENCES DE LA TERRE