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Post-conference workshop WS07: Greenstone Belt Architecture and Metal Endowment of the Superior Craton

Tectonic evolution of the western Wabigoon terrane in the Sturgeon Lake region: Implications for understanding structural controls on the greenstone belt and its VMS deposits

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Outline

PART I

Crustal architecture of the western Wabigoon terrane in the Sturgeon Lake region

PART II

<u>New stratigraphic, geochronologic, and isotopic data</u> from the Sturgeon Lake greenstone belt

PART III

<u>Tectonic evolution</u> of the western Wabigoon terrane in the Sturgeon Lake region and <u>structural control</u> on VMS deposits

PART I

Crustal architecture of the western Wabigoon terrane in the Sturgeon Lake region









Geologic background

MEST = Metal Earth Sturgeon transect WS1a & WS1f = Lithoprobe transects



Geologic background

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Western Wabigoon terrane (WWT)

- 4 volcanic assemblages (~2775-2720 Ma)
- Syn-volcanic & post-tectonic intrusions
- VMS deposits (Zn-Cu-Pb-Au-Ag)
- Folded at the map scale (syncline)
- Two major fault zones

Winnipeg River terrane (WRT)

 Basement rocks intruded by syn- to posttectonic granitoids with a variety of textures and compositions.

Major deformation between WWT and WRT:

~2700–2695 Ma (Sanborn-Barrie and Skulski, 2006)

Map modified after Sanborn-Barrie and Skulski (2005) and Stone et al. (2002)



Geologic background

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Cross section along the Sturgeon transect



Reflection seismic survey

Pre-stack migrated seismic reflection of Sturgeon transect

70 km long & 48 km deep

Source spacing: 50 m

Receiver spacing: 25 m



Reflection seismic survey

- 7 zones of distinctive reflectivity
- 2 fault zones



Reflection seismic survey

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Zones B, C, D, E, F:

Prominent reflections gneissic fabrics and/or mafic layers

Poorly reflective zones - granitoid layers and bodies

= Winnipeg River terrane (gneisses and granitoids intruded by younger granitoids)

The south-dipping reflectors in Zone C: mafic-ultramafic dikes imaged by out-of-plane seismic ray (exclusively above the apparent crustal root)



Zone G: Minimum thickness of Sturgeon Lake greenstone belt

Greenstone basal boundary:

- Beneath Zone G & above the Zone F reflectors
- A thrust fault zone

An apparent crustal root - the middle part of Moho with deeper depths; possibly caused by partial melting of mantle rocks that would result in interlayered felsic and mafic rocks, similar in reflectivity to that of the lower crust 10 S 12 11 10 (A) Magnetotelluric survey 20 **Cross-section of 3D** 25 inversion model 30 along the seismic line 40 45 Distance (Km Winnipeg River terran (B) South Sturgeon Distance (Km) 25 0 Warm colors: **R**2 5 conductive zone 10 CT 15 Cold colors: (Km) 20 resistive zone 25 30 35 Reflection Moho 40

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The Smith

Horizontal slices of resistivity structures from surface to mantle

20°W -91.04°W





Upper-crustal conductive fingers: Interconnected sulfides and/or high iron content in Fe/Mg-bearing silicates

- Interconnected grain-boundary graphite films
- Diffusion of hydrogen in nominally anhydrous silicates due to extensive crustal partial melting involved mantle components

alic pluton

Interpretation:

Crustal structure beneath Sturgeon transect



Lithoprobe seismic transects





Lithoprobe WS1f line (E-W)



Ma et al. (2021)

Composite crustal section (South-North) (Sturgeon transect + partial WS1a)



• Continuous distribution of greenstone belt as a 5–10 km thick thrust sheet in the upper crust

Ma et al. (2021)

- A collision zone in the mid- to lower-crust
- A trans-crustal fault zone, F1, accommodated the collision of basement and the emplacement of greenstone belt
- Subcreted crust in the mantle

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• Thickened crust beneath the greenstone belt

3D crustal structure

Sturgeon transect + Lithoprobe WS1a & WS1f

Miniss

River

Sioux Lookout

-ME

Fig. 2

PART II

New stratigraphic, geochronologic, and isotopic data from the Sturgeon Lake greenstone belt

¹⁷ Stratigraphy of Upper Handy Lake assemblage

Upper

²⁰ Stratigraphy of Upper Handy Lake assemblage

Upper

Handy

Lake

assemblage

Fsp- & qtz-phyric felsic flow

Fsp-phyric mafic flow

Qtz- & fsp-phyric felsic crystal tuff

Interbedded felsic tuff & felsic lapilli tuff

(Lapilli) tuff

Flow

Pyroclastic breccia

Felsic & mafic blocks supported by intermediate lapilli tuff

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Felsic blocks supported by qtzphyric felsic crystal tuff

New geochronological and isotopic results (TIMS)

Felsic flow of Upper Handy Lake assemblage

2734 ± 1 Ma

Published zircon ages from the VMS-bearing South Sturgeon assemblage:

Mattabi rhyolite	>2734.7 ± 1.6
NBU rhyolite	2735 ± 1.7
NBU rhyolite	2735.5 + 3/-1.9
NBU rhyolite	2736.3 +9.3/-3.9
Lyon Lake rhyolite	2735.2 + 6.9/-3.2
Lyon Lake rhyolite	2734.8 + 2.8/-2.5
Lyon Lake rhyolite	2736 ± 1.8

Davis et al. (1985)

New geochronological and isotopic results (LA-ICPMS)

4 synvolcanic granitoids

- South Sturgeon assemblage (n=2)
- Shanty Lake pluton (n=1)
- Lower Handy Lake granite (n=1)

New geochronological and isotopic results (LA-ICPMS)

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Synvolcanic granitoid – in contact and coeval with the lowermost South Sturgeon assemblage

New geochronological and isotopic results (LA-ICPMS)

New geochronological and isotopic results (LA-ICPMS) Zircon depleted mantle Hf model ages of synvolcanic granitoids

PART III (conclusions)

Tectonic evolution of the western Wabigoon terrane in the Sturgeon Lake region: Implications for understanding structural controls on the greenstone belt and its VMS deposits

- A *collision belt* between the main WRT to the north and a continental margin promontory of the WRT to the south.
- Prior to the collision, the WWT was partially *accreted onto* and partially *subducted* under the WRT.

- Prior to the accretion, various volcanic and plutonic rocks of WWT were developed in an oceanic crust setting from ~2775 to ~2720 Ma.
- The felsic intrusive rocks with *inherited zircons* of 2880, 2820, & 2810 Ma are *juvenile* and were melted from older juvenile sources that were *episodically* extracted from the mantle likely in the Mesoarchean.

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South Sturgeon VMS deposits

- In this model, the South Sturgeon VMS deposits were tectonically emplaced to the current position, implying that thrusting and folding of the greenstone belt did not liberate the VMS deposits.
- The newly identified 2734±1 Ma felsic flow in the Upper Handy Lake assemblage demonstrates occurrence of age-equivalent rocks to the VMSbearing South Sturgeon assemblage in the north limb of the regional syncline.
- The stratigraphy of the upper part of the UHLA is consistent with that of the South Sturgeon felsic volcanic rocks.
- Prospective for VMS in this new ~2734 Ma unit (?)

The Beidelman Bay pluton is considered to have initiated and sustained the sub-seafloor convective system responsible for the formation of the South Sturgeon VMS deposits (Galley et al., 2000).

The *newly discovered 2738.6 ± 1.3 Ma felsic intrusion* at the base of the South Sturgeon assemblage may have jointly facilitated the formation of the VMS deposits.

Neoarchean trans-crustal fault zone

• The Neoarchean trans-crustal fault zone F1 extends from surface into the mantle, which may have implications for gold prospecting, as crustal-scale fault zones are commonly related to orogenic gold (e.g., Frieman and Kuiper, 2019).

Project Research & Funding Partners

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Thank you!

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