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NATURAL RESOURCES CANADA - INVENTIVE BY NATURE


Geology, Genesis and Exploration for Magmatic Cr and Fe-Ti-V Ore Systems

***PDAC Short Course – Saturday March 2nd, 2019
Geology, Genesis, and Exploration for Magmatic
and Magmatic-Hydrothermal Ore Deposits***




TGI
Ore Systems


Michel G. Houlé
Geological Survey of Canada – Québec




Laurentian University
Université Laurentienne




HARQUAIL SCHOOL OF EARTH SCIENCES
ÉCOLE DES SCIENCES DE LA TERRE



MERC
Mineral Exploration Research Centre
AT THE HARQUAIL SCHOOL OF EARTH SCIENCES



METAL EARTH



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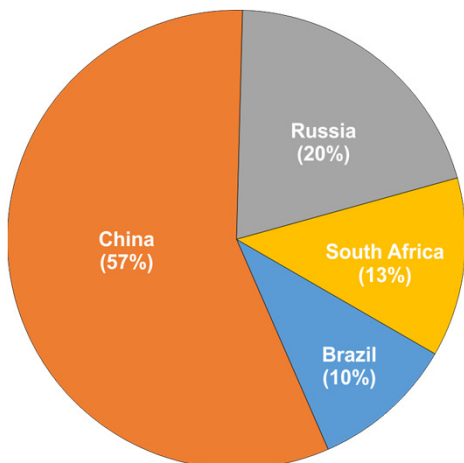


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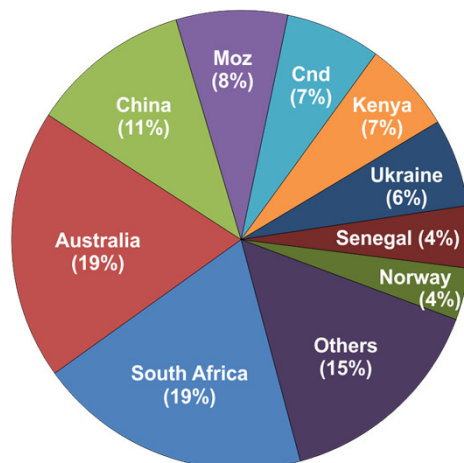
3

World V and Ti Production in 2016-V / 2017-Ti

Vanadium Production (2016)



Titanium Production (2017)



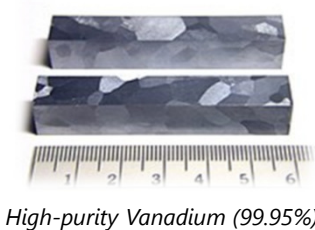
Data from the USGS Mineral Resources Program

4

Uses and Applications

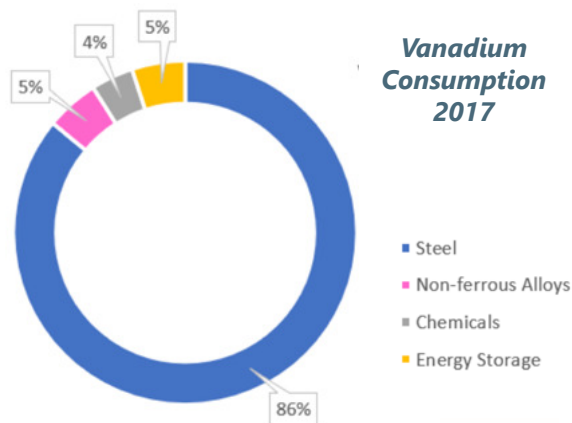
Properties

- ❖ Vanadium is a medium-hard, ductile, bright steel-blue metal
- ❖ Has good resistance to corrosion and it is stable against alkalis and sulfuric and hydrochloric acids



Uses

- ❖ Most V is used as steel additive (ferrovanadium) steel to make it stronger
- ❖ High-strength low-alloy containing V are widely used for auto parts, buildings, bridges, etc.
- ❖ Aerospace application: V-Ti alloys have best strength-to-weight ratio
- ❖ Is becoming more widely used in green technology (Vanadium Redox Flow Batteries)



Vanadium Consumption 2017

Schulz et al. 2017 – USGS PP1802

5

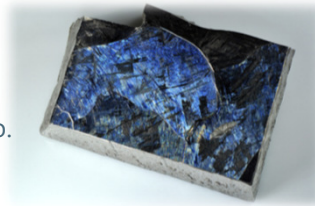
Uses and Applications

Properties

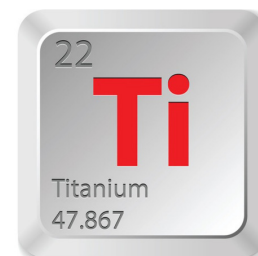
- ❖ Titanium is recognized for its high strength-to-weight ratio.
- ❖ It is a strong metal with low density that is quite ductile, lustrous, and metallic-white in color

Uses

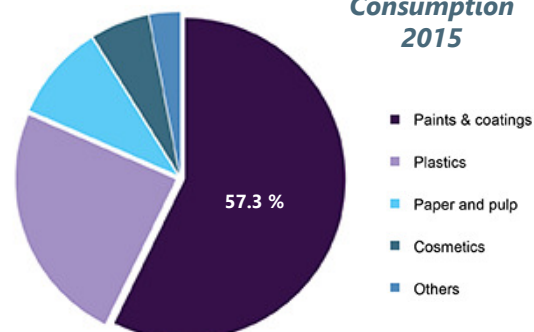
- ❖ Pigments, additives, and coatings (paints, paper, toothpaste, and plastics)
- ❖ Aerospace and marine (as steel additive)
- ❖ Industrial (pipe and process equipment such as heat exchangers, tanks, etc.)
- ❖ Consumer and architectural (automotive applications, sporting goods, etc.)
- ❖ Jewellery (Titanium ring, etc.)
- ❖ Medical (surgical implements and implants, such as hip balls, joint replacement and dental implants)



High-purity (99.999%) titanium
with visible crystallites



Titanium Consumption 2015



Source: Grand View Research

6

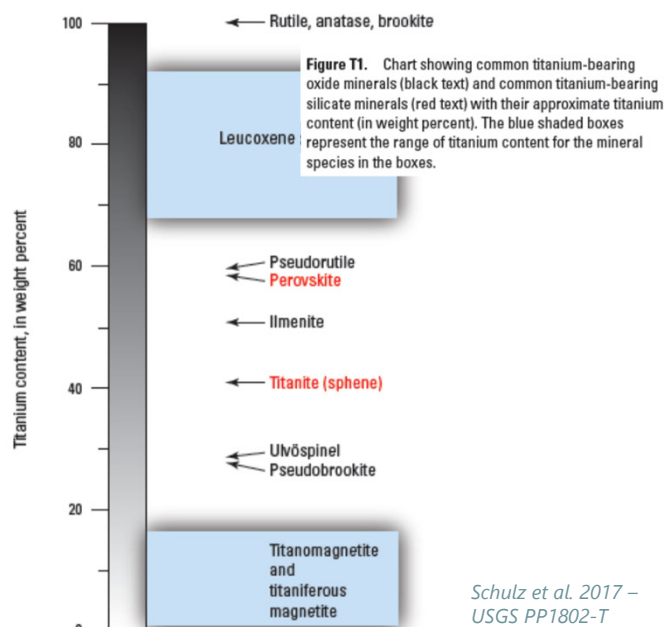
Mineralogy

Vanadium-bearing minerals

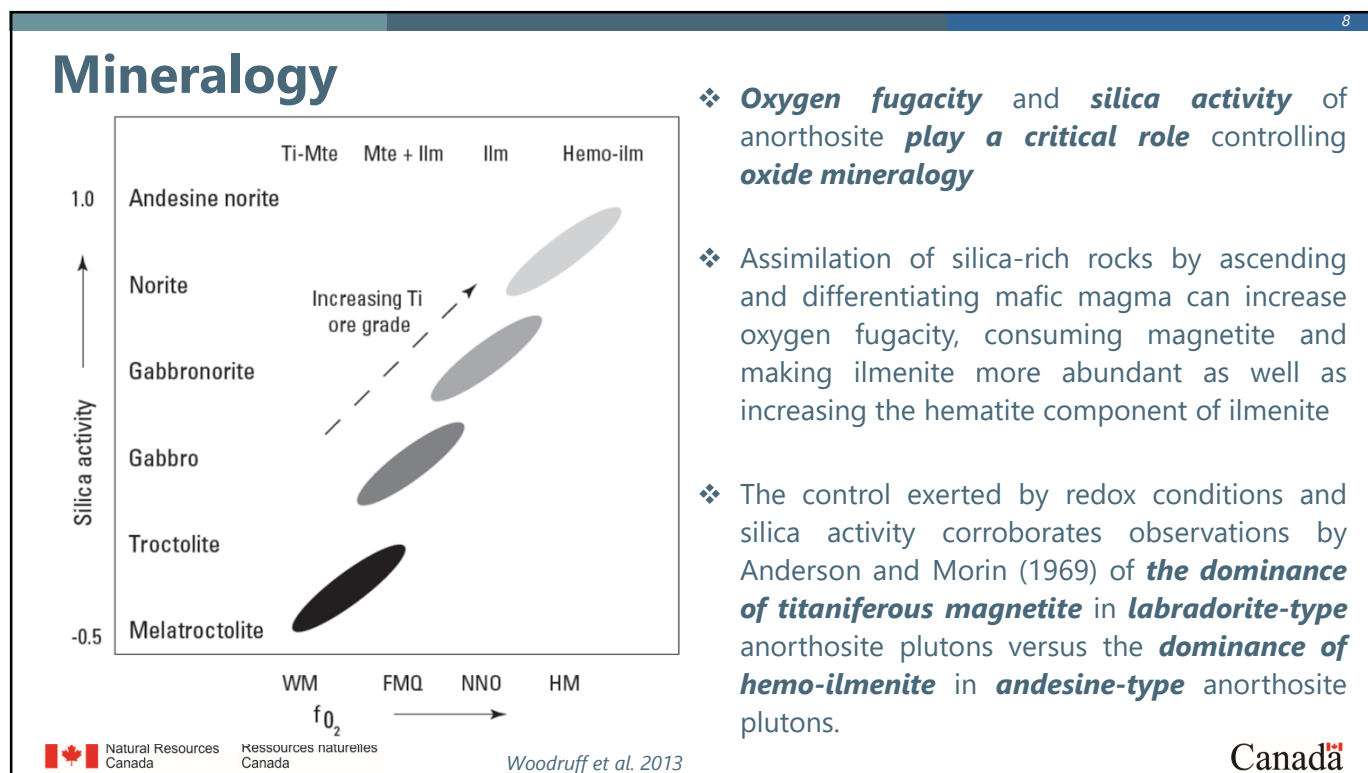
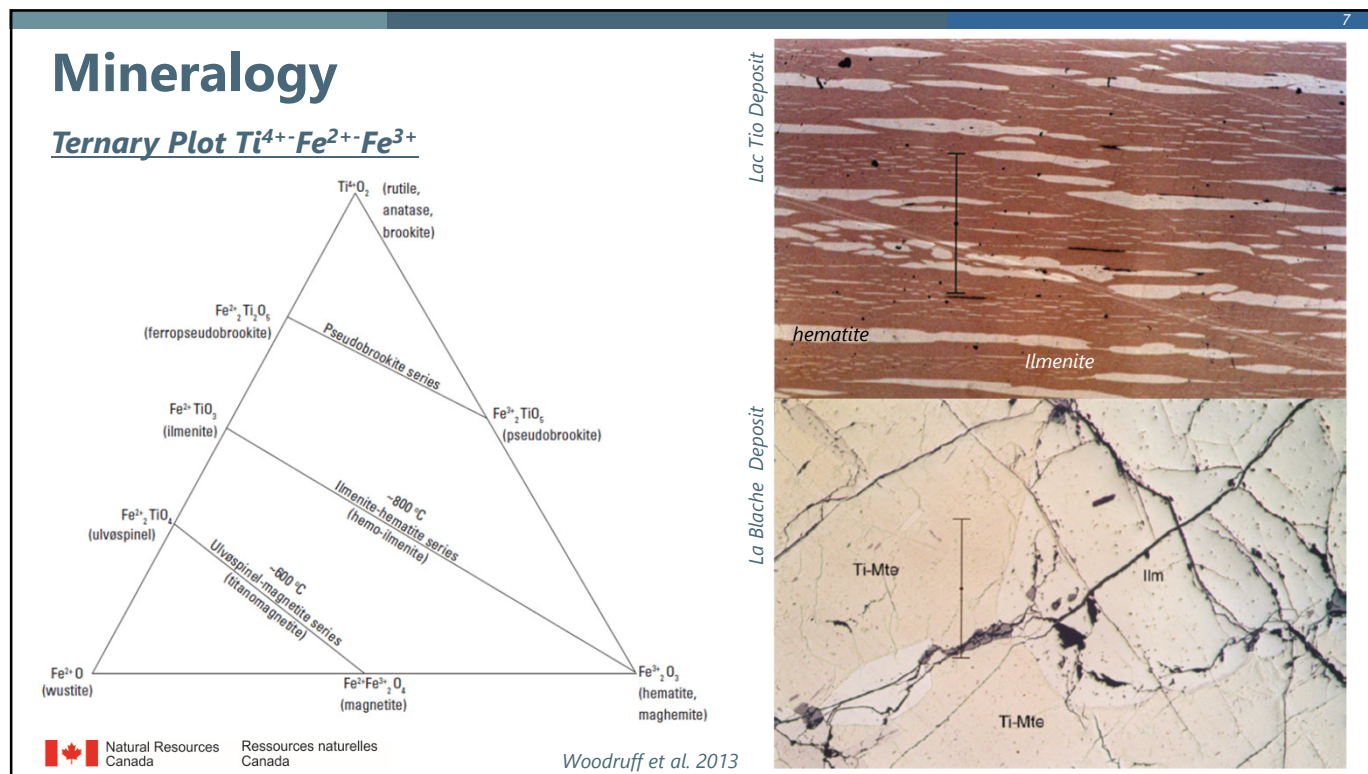
- ❖ Coulsonite – $\text{Fe}^{2+}\text{V}^{3+}_2\text{O}_4$
- ❖ Hematite – Fe_2O_3
 - ❖ Common impurities: Ti, Al, Mn, H₂O
- ❖ Ilmenite – $\text{Fe}^{2+}\text{TiO}_3$
 - ❖ Common impurities: Mn, Mg, V
- ❖ Magnetite – $\text{Fe}^{2+}\text{Fe}^{3+}_2\text{O}_4$
 - ❖ Common impurities: Mg, Zn, Mn, Ni, Cr, Ti, V, Al
- ❖ Perovskite – CaTiO_3
 - ❖ Common impurities: Fe, Nb, Ce, La, TR
- ❖ Rutile – TiO_2
 - ❖ Common impurities: Fe, Ta, Nb, Cr, V, Sn

Schulz et al. 2017 – USGS PP1802-U

Titanium-bearing minerals



Schulz et al. 2017 –
USGS PP1802-T



Classification Scheme for V and Ti Deposits

Vanadium

- ❖ **Vanadiferous Titanomagnetite Deposits**
 - ❖ *Magnetite-ilmenite dominant minerals*
 - ❖ *Hematite, perovskite, rutile present in some deposits*
- ❖ Sandstone-Hosted Vanadium Deposits
- ❖ Shale-Hosted Vanadium Deposits
- ❖ Vanadate Deposits
- ❖ Other Magmatic-Hydrothermal Vanadium Resources
- ❖ Fossil Fuels

Titanium

- ❖ **Igneous Deposits**
 - ❖ *Massif anorthosite – Hemo-ilmenite (very high)*
 - ❖ *Anorthosite-gabbro – Ilmenite, titanomagnetite (high)*
 - ❖ *Layered mafic intrusion – Ilmenite, titanomagnetite (high)*
 - ❖ *Troctolite/ultramafic – Ilmenite, titanomagnetite*
 - ❖ *Albitic/metasomatized – Rutile*
 - ❖ *Alkalic/metasomatized – Perovskite, brookite, rutile*
 - ❖ *Weathered alkali rocks – Anatase*
 - ❖ *Weathered anorthositic rocks – Ilmenite, rutile*
- ❖ Metamorphic Deposits
- ❖ Sedimentary/heavy mineral Deposits

Schulz et al. 2017 – USGS PP1802-U



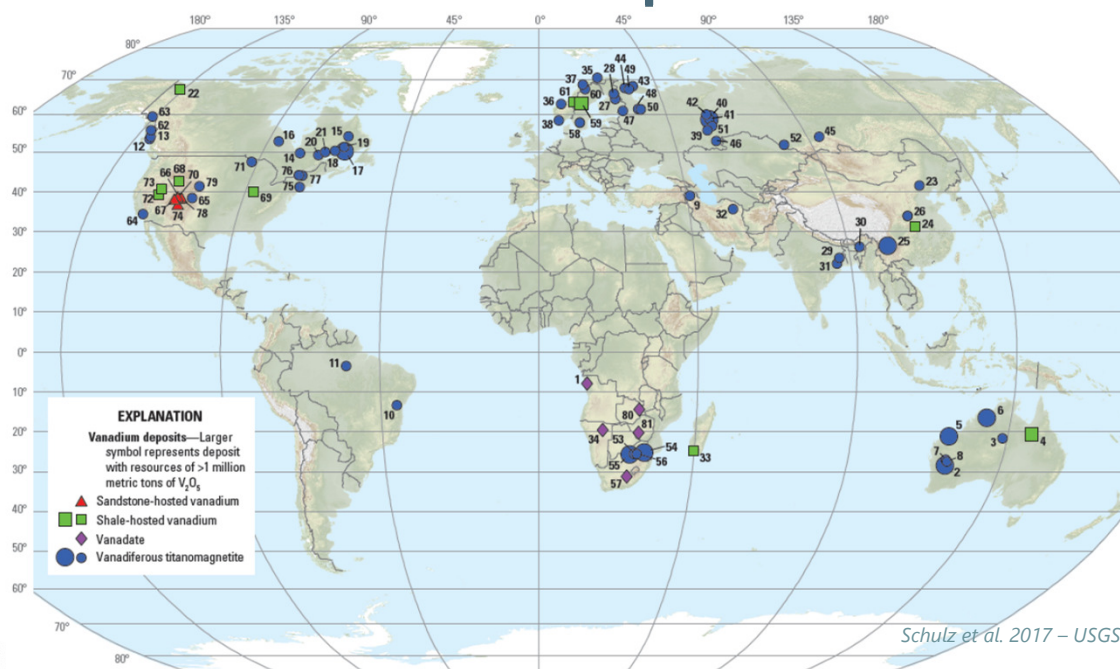
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Schulz et al. 2017 – USGS PP1802-T

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Distribution of Vanadium Deposits



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Schulz et al. 2017 – USGS PP1802-U

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11

Distribution of Titanium Deposits



12

Classification Scheme for Fe-Ti-V-(P) Deposits

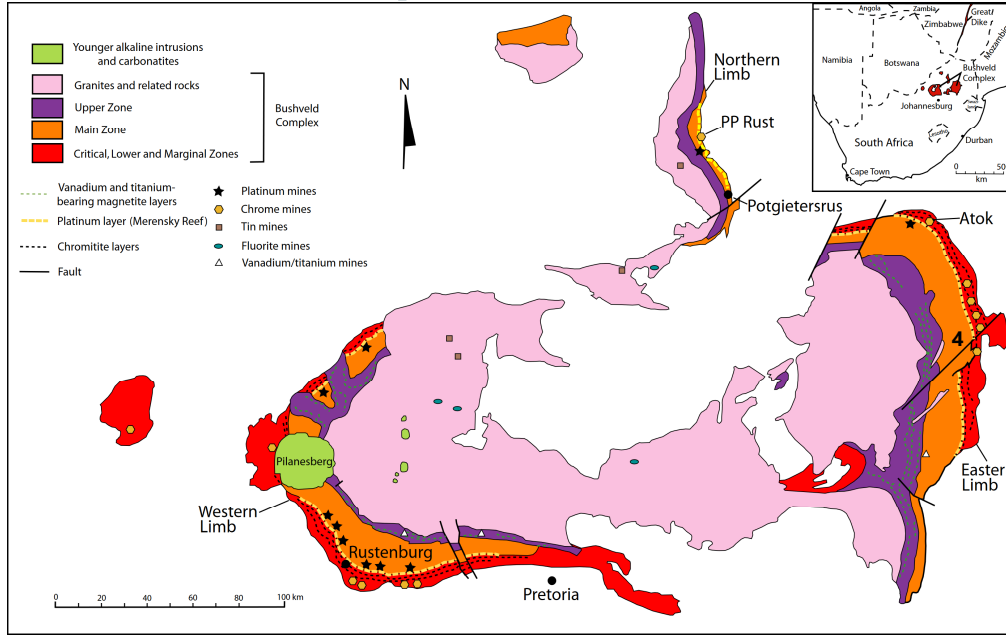
Enrichment of magnetite, titanomagnetite, and ilmenite may occur in both ultramafic and mafic rock types and several distinct types of deposits occurred and may be classified in many ways

- ❖ Upper Parts of Large Layered Complexes
 - ❖ Examples: Bushveld Complex (South Africa), Windimurra Complex (Australia), Lac Doré Complex-? (Canada)
- ❖ Flood basalt-related/Greenstone belt-related deposits
 - ❖ Panzhihua, Hongge, Baima, Taihe (Emeishan LIP – China), Ring of Fire prospects-? (Canada)
- ❖ Anorthosite-related deposits: Proterozoic Anorthosite Complexes
 - ❖ Examples: Lac Tio (Canada), Tellnes (Norway), Damiao (China), St-Urbain (Canada)
- ❖ Alaskan-Urals Type-related deposits
 - ❖ Examples: Kachkanar (Russia)

Modified from Pang et al. 2010 - Lithos

13

Bushveld Complex – South Africa



World Mineral Treasure: Bushveld Igneous Complex

- ❖ World's largest layered mafic-ultramafic complex
- ❖ Covering 66,000 km²
- ❖ World largest chromite and PGEs resources
- ❖ Chromite is confined within the Critical Zone
- ❖ PGE is confined at the top of the Critical Zone near the Main Zone
- ❖ V mineralization occur throughout the Upper Zone
- ❖ Continuous over 100s km



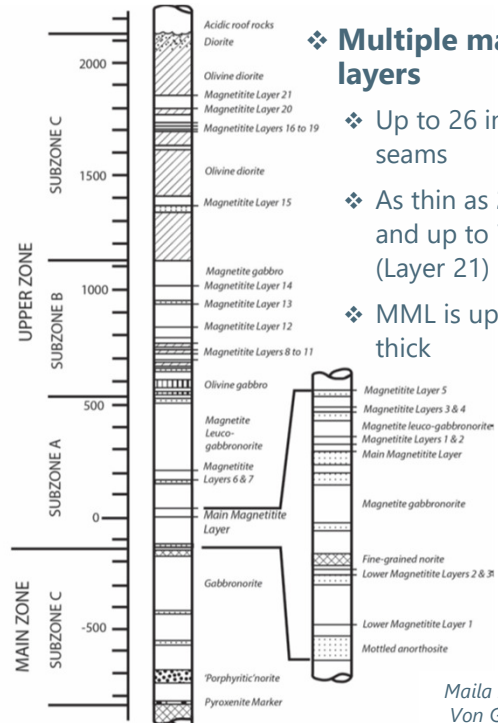
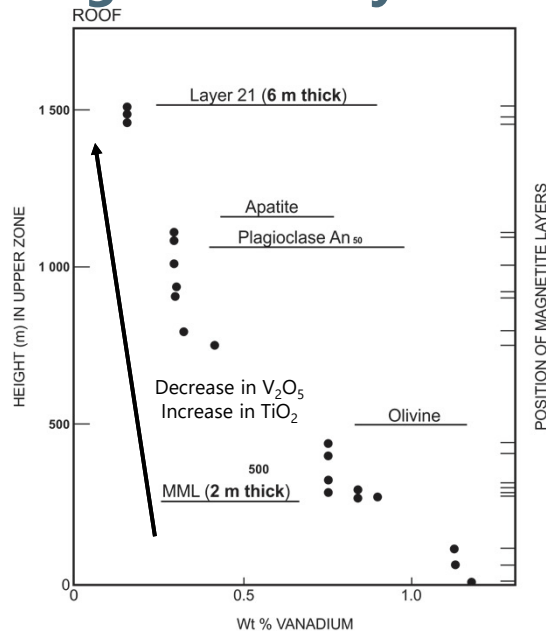
Natural Resources Canada / Ressources naturelles Canada

Taylor et al. 2009 – USGS

Canada

14

Magnetitite Layers



Multiple magnetite layers

- ❖ Up to 26 individual seams
- ❖ As thin as 2 cm-thick and up to 7.1 m thick (Layer 21)
- ❖ MML is up to 2.5 m thick



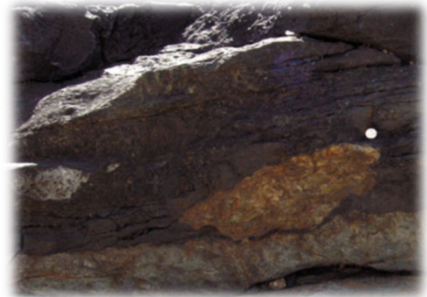
Natural Resources Canada / Ressources naturelles Canada

Cawthorn 2015 – CGSA

Maila 2015; Harney and Von Gruenewaldt, 1995

15

Main Magnetite Layer



Typical textural facies within the Main Magnetite Layer – Magnet Heights



Natural Resources Canada

Scoon & Mitchell - 2012
Ressources naturelles Canada

Photos (central and right ones) from C. Rowan - 2008

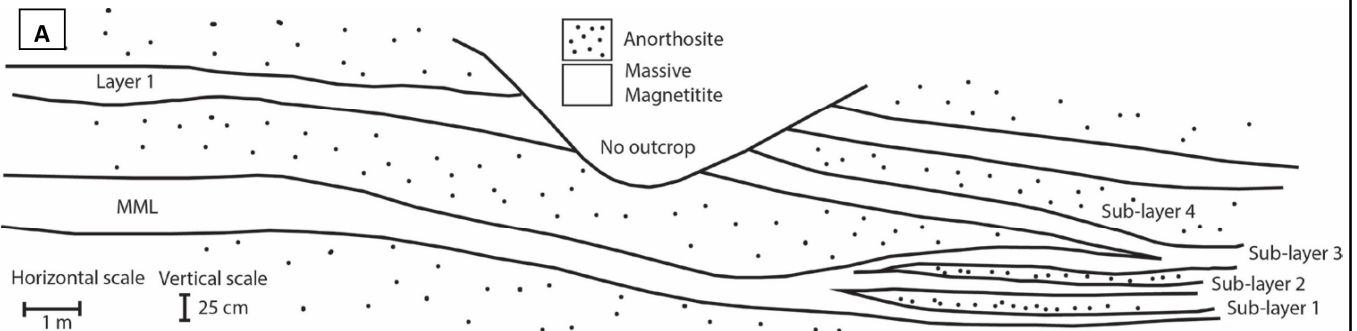
16

Magnetitite Layers



❖ Magnetitite bifurcations

- ❖ Sketch of MML spitting into three sub-layers and merging with layer 1
- ❖ Small scale magnetite bifurcation in samples

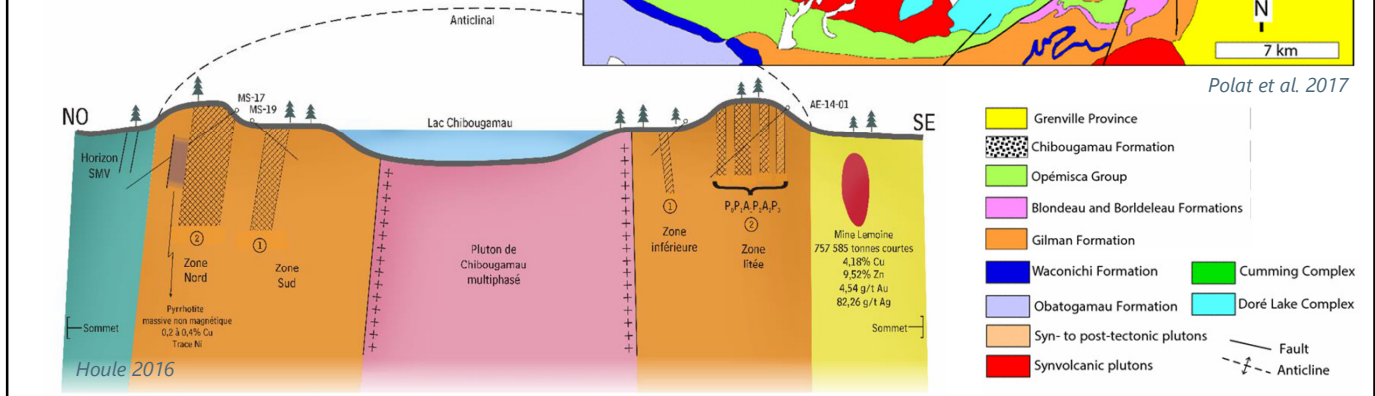


Maila 2015 - Thesis

Lac Doré Complex

❖ Large layered intrusion

- ❖ Fe-Ti-V mineralization occur on both limbs of the anticline
- ❖ However, the south limb is where the best mineralization is found and contain several deposits



Lac Doré Complex - Stratigraphy

Flanc sud du complexe			
	ZONES	SOUS-ZONES	LITHOLOGIE
SÉRIE SUPÉRIEURE	BORDURE		Gabbro Anorthosite gabbroïque
	GRANOPHYRE		
SÉRIE LITÉE	ZONES LITÉES	P ₃	À quartz À apatite et ilménite À ilménite et ferrosilicates Ferropyrroxénite Ferrogabbro
		A ₂	Anorthosite gabbroïque
		P ₂	Magnétite Ferropyrroxénite Ferrogabbro
		A ₁	Anorthosite gabbroïque
		P ₁	Magnétite Ferropyrroxénite Ferrogabbro
SÉRIE INFÉRIEURE	ANORTHOSIQUE s.l.		Gabbro Gabbro anorthosique Anorthosite noritique Anorthosite (s.s.)
	P ₀	Magnétitique	Ferrogabbro Ferropyrroxénite à magnétite

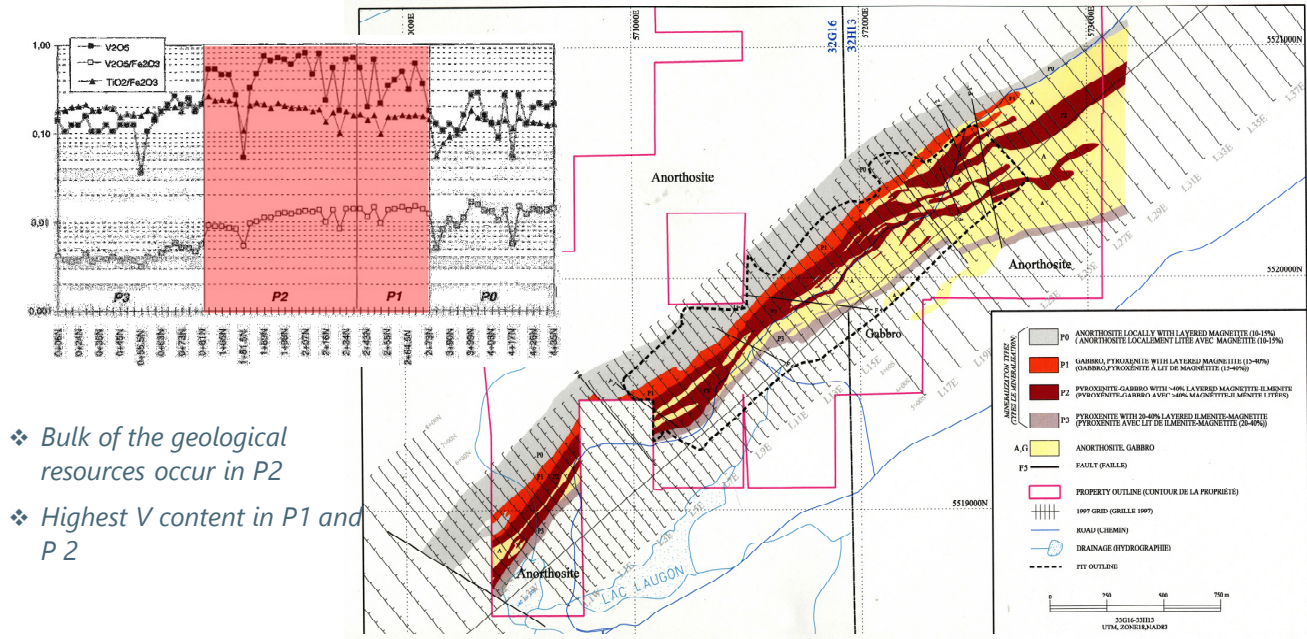
- ❖ **P0:** Anorthositic sequence locally with layered magnetite (10-15%)
- ❖ **P1:** Layered sequence of anorthosite, gabbro and pyroxenite interlayered with magnetite layers (15-40%). Highest V content
- ❖ **P2:** Pyroxenitic and gabbroic sequence with abundant interlayering of massive to semi-massive magnetite-ilmenite layers (>40%) and magnetite/ilmenite-bearing ferrogabbro.
- ❖ **P3:** Magnetite/ilmenite bearing ferrogabbroic sequence interlayered with magnetite/ilmenite layers (20-40%)

	Thickness as calculated from sampling	V ₂ O ₅	Fe ₂ O ₃	TiO ₂	% Mgt from visual estimation
P0	32 m	0.194%	20.1%	1.74%	10%
P1	49 m	0.340%	41.5%	4.27%	15-30%
P2	97 m	0.486%	9.20%	9.20%	>40%
P3	38 m	0.163%	48.6%	6.43%	20-30%
Total	216 m	0.353%	49.8%	6.49%	

Daigneault et Allard, 1990; Girard and Allard 1998; Girard and D'Amours 2015

19

Lac Doré Complex – Ore Zone



- ❖ Bulk of the geological resources occur in P2
- ❖ Highest V content in P1 and P2



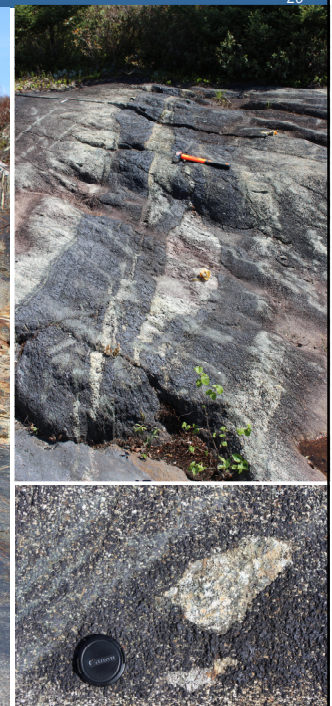
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Girard and Allard 1998; Girard and D'Amours 2015

20

Typical Fe-Ti-V Mineralization in P₂



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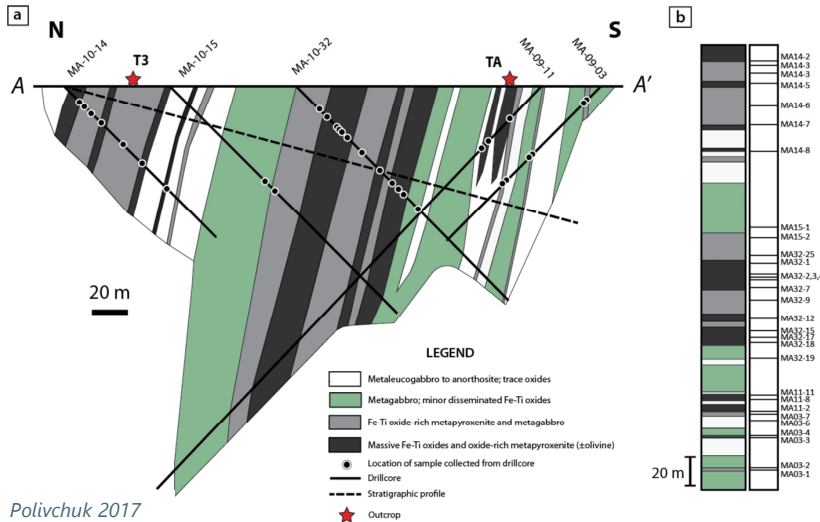
Ressources naturelles Canada

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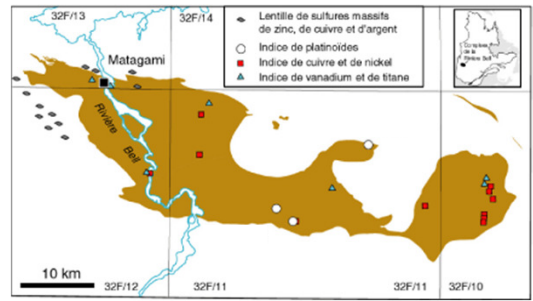
21

Rivière Bell Complex

Iron-T deposit is hosted by magnetite-ilmenite pyroxenites and gabbros



Polivchuk 2017



Goutier 2015

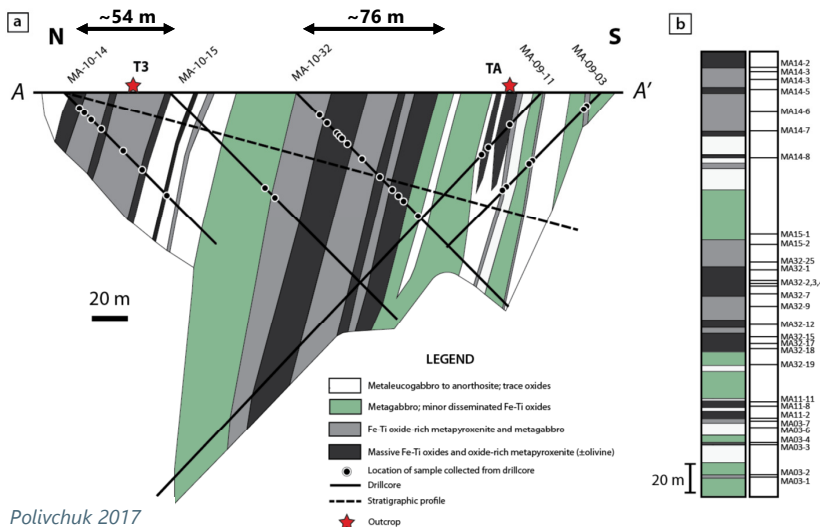
Iron-T deposit is hosted by magnetite-ilmenite pyroxenites and gabbros

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22

Rivière Bell Complex

- ❖ The oxide-rich gabbro horizons varying in width from 10 to 100 m clearly appear on the airborne regional magnetic survey
- ❖ *Iron-T deposit is hosted by magnetite-ilmenite pyroxenites and gabbros*



Polivchuk 2017



Goutier 2015

- ❖ The oxide-rich gabbro is a mineralized cumulate forming either homogeneous horizons with disseminated oxide mineral contents ranging from 20 to 60% or homogeneous massive layers with oxide mineral contents varying from 60 to 90%
- ❖ Massive oxide mineralized bands are interlayered with poorly mineralized gabbro forming pluri-centimetric to decimetric scale interlayers

Taner et al. 1998

Canada

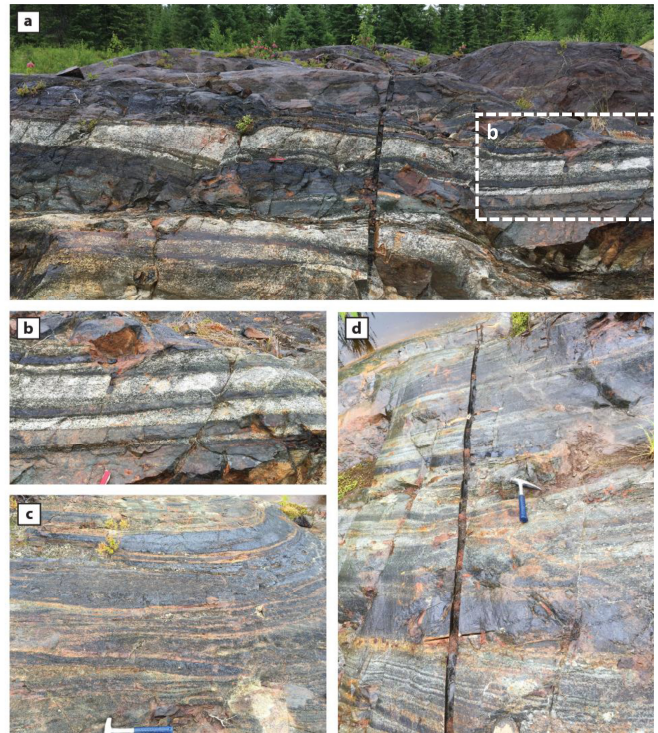
23

Rivière Bell Complex

Igneous layering textures of the RBC within the Iron-T property at outcrop T3

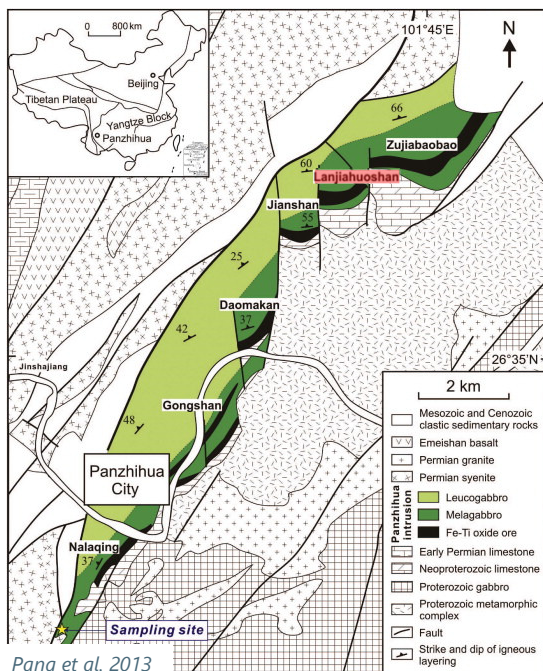
- Interlayered centimeter-to-decimeters thick, variably oxide-rich metaleucogabbro (speckled white layers) and metapyroxenite (dark green-grey layers). Swiss army knife for scale
- Close-up of a sequence of several centimeters-thick oxide-rich pyroxenites and leucogabbros. Melanocratic bands exhibit both sharp and modally graded upper contacts with overlying leucogabbros. Note the tapering lowermost leucocratic band
- Discontinuous ribbon-like banded massive oxides and gabbros. Massive oxides exhibit sharp upper and lower contacts
- Cyclically layered magnetite gabbro. Some layers exhibit internal normal grading of oxides and mafic minerals.

Polivchuk 2017 ices Ressources naturelles Canada



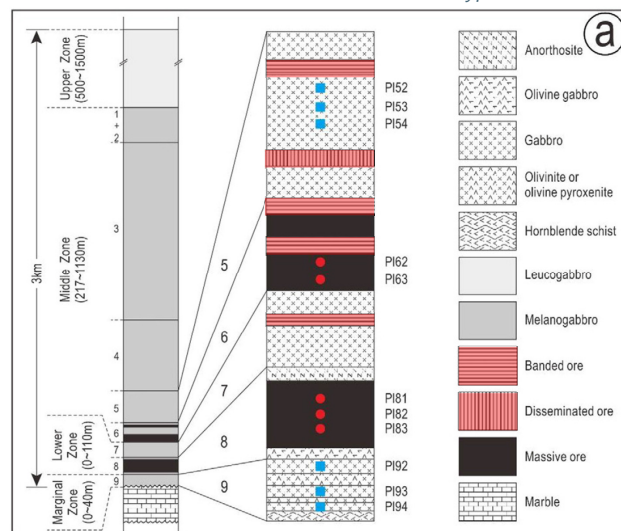
24

Panzhihua Intrusion



From the base upwards, the Panzhihua intrusion is divided into four zones based on textural features and cumulus minerals

- ❖ Marginal Zone; Lower Zone; Middle Zone; and Upper Zone
- ❖ Subdivided into 9 orebodies based on mineralization type

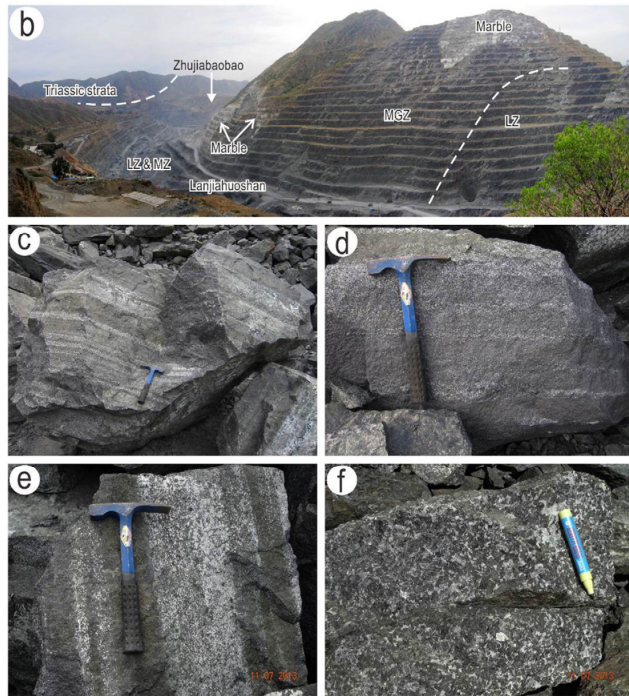
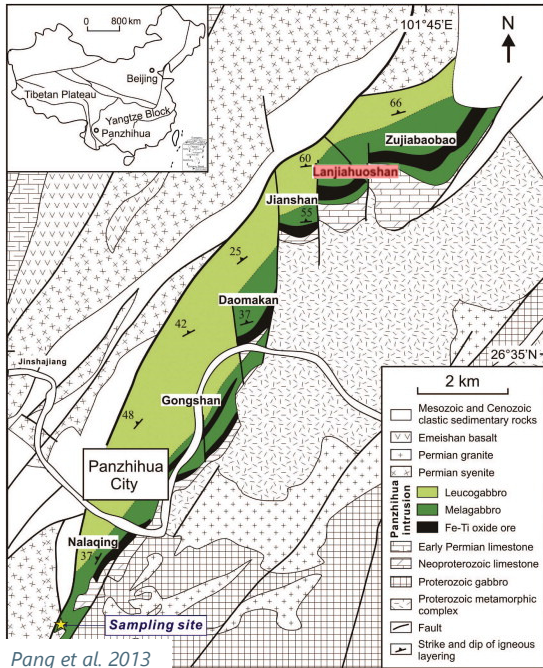


Gao et al. 2017

Canada

25

Panzhihua Intrusion

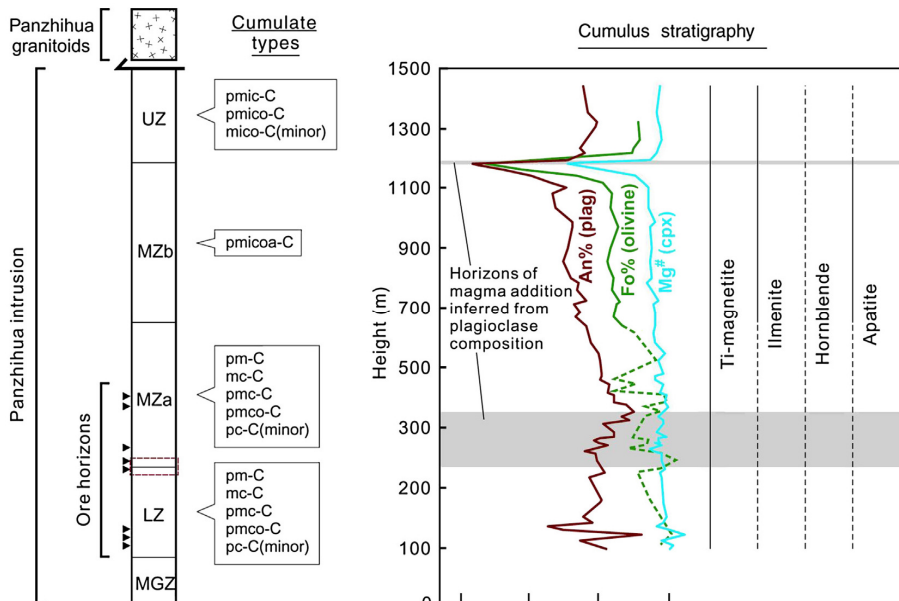


Gao et al. 2017

Canada

26

Panzhihua Intrusion



Simplified stratigraphic column of the Panzhihua intrusion

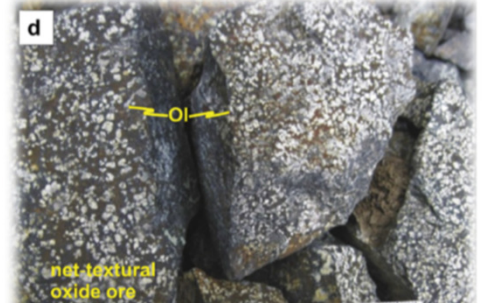
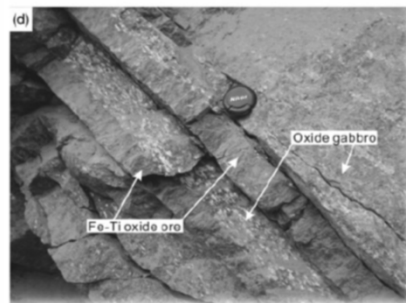
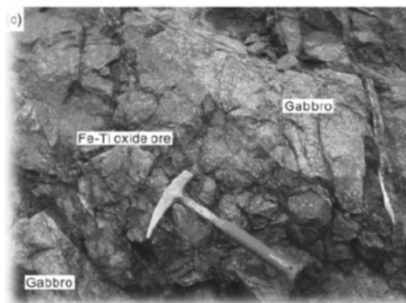
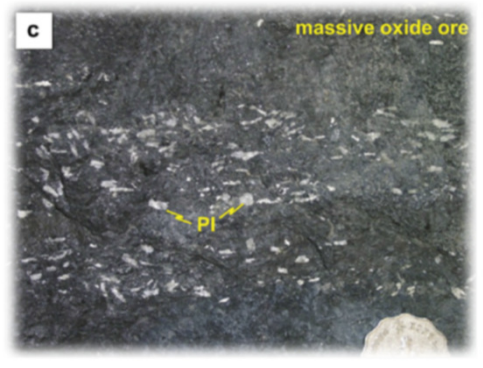
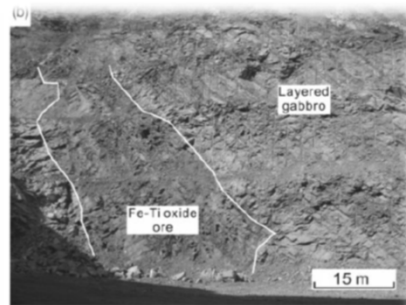
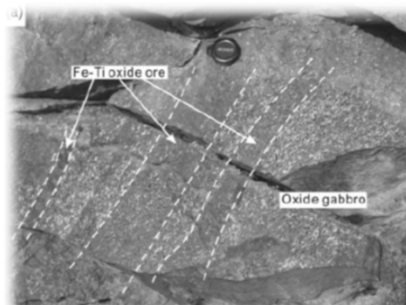
- ❖ showing the zonal division and the variations in cumulate types, cumulus minerals and their compositions.
- ❖ The cumulate nomenclature
 - ❖ "-C" meaning cumulus:
 - ❖ Plagioclase (p)
 - ❖ magnetite (m)
 - ❖ ilmenite (i)
 - ❖ clinopyroxene (c)
 - ❖ olivine (o)
 - ❖ apatite (a).

Pang et al. 2013

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27

Panzhihua Intrusion

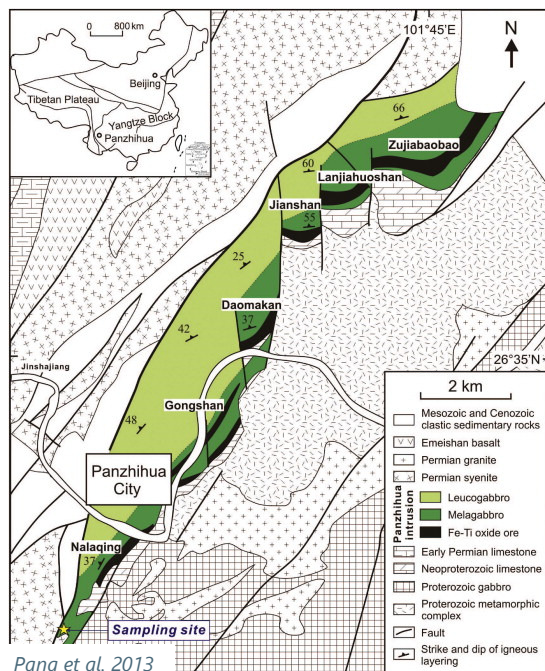


Pang et al. 2008

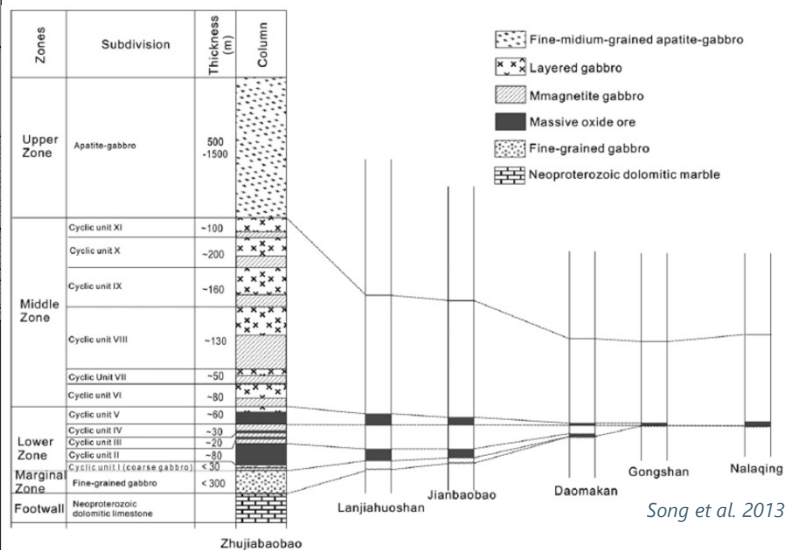
Zhou et al. 2013

28

Panzhihua Intrusion



Pang et al. 2013

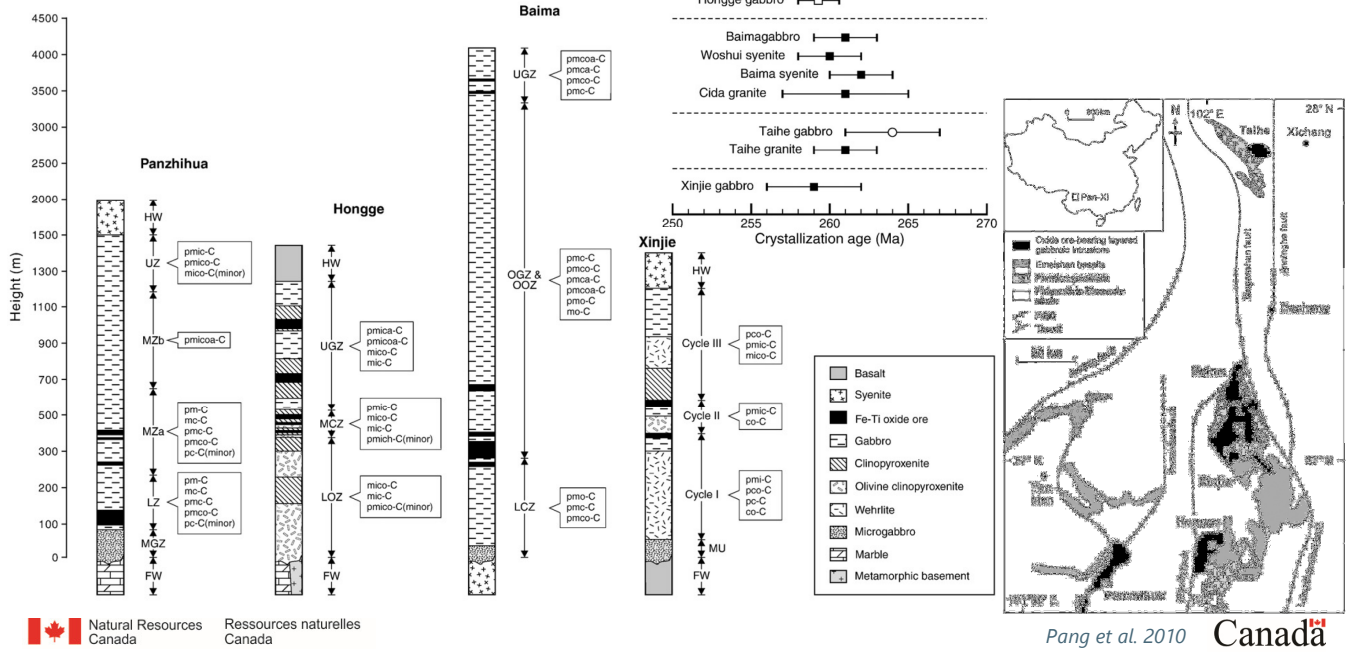


Song et al. 2013

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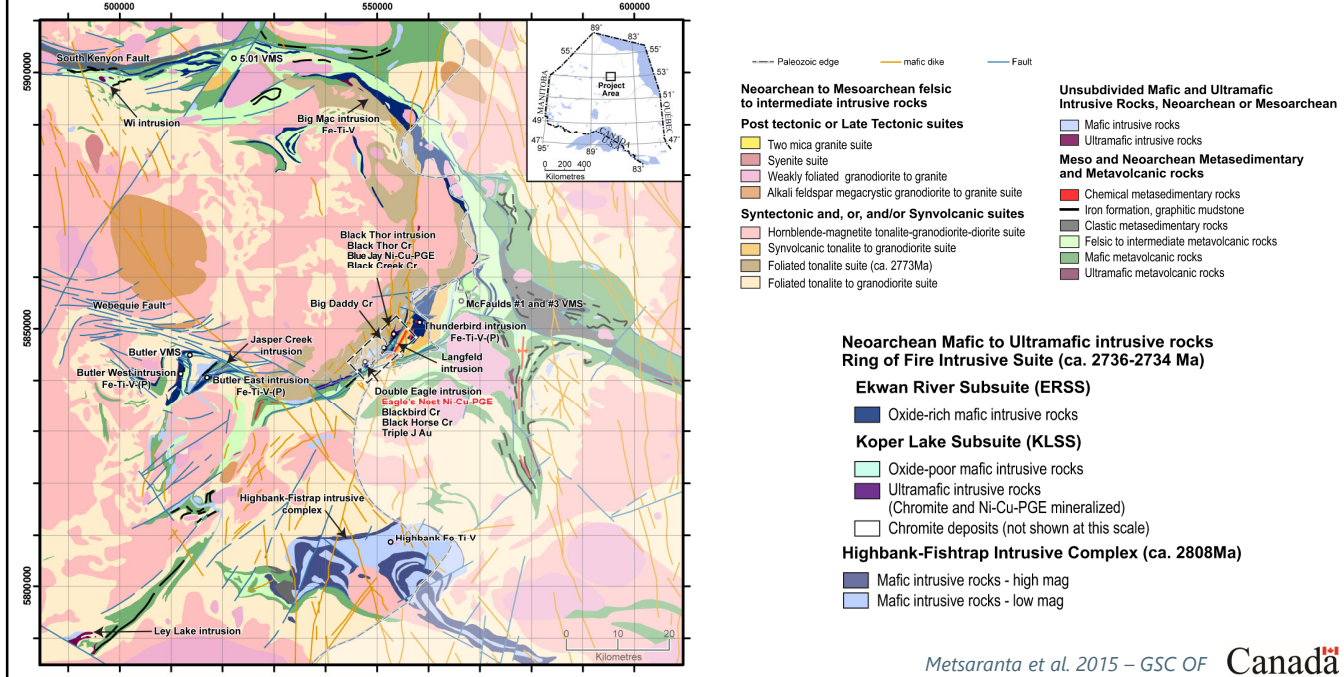
29

Panzhihua-Type Bodies

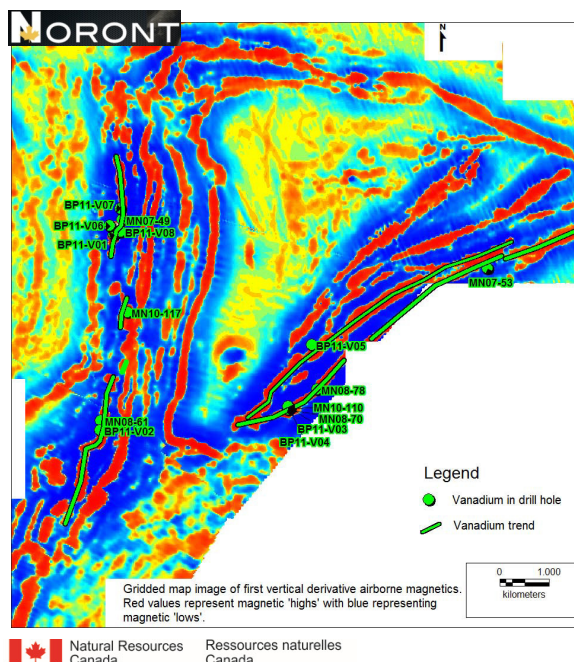


30

Fe-Ti-V-(P) Mineralization - RoF

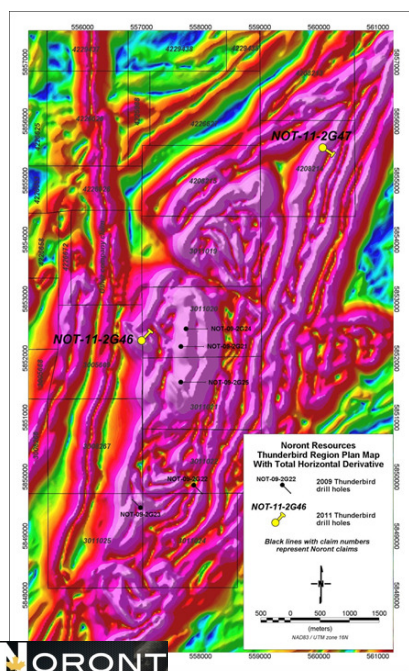


Fe-Ti-V-(P) Mineralization – Butler Area



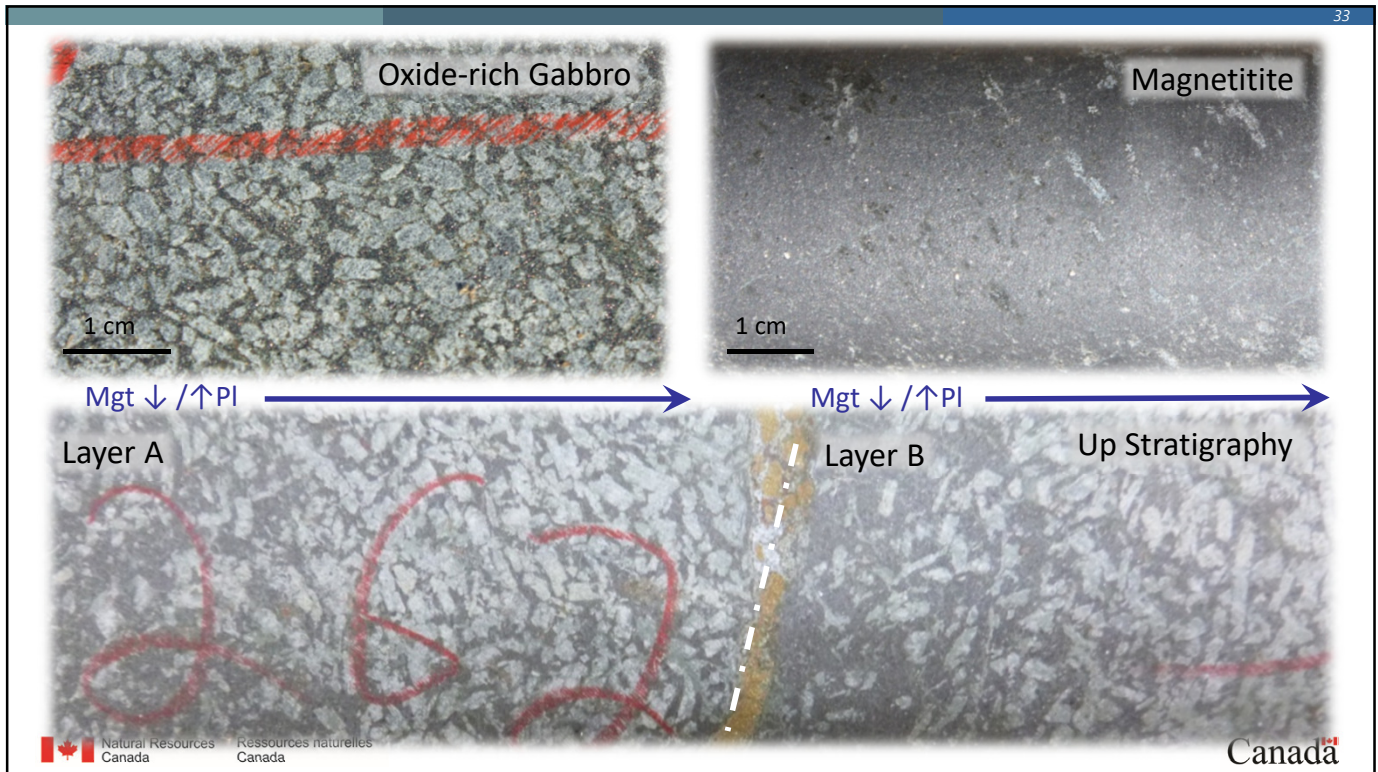
- ❖ Several Fe-Ti-V mineralized horizons
 - ❖ Both flanks (Western & Eastern)
- ❖ Fe-Ti-V mineralization
 - ❖ 1.17% V₂O₅, 7.97% TiO₂ and 40.73% Fe over 4.2 m as higher grade intersection
 - ❖ 0.54% V₂O₅, 0.46% TiO₂ and 23.78% Fe over 36 m as a wider, lower grade

Fe-Ti-V-(P) Mineralization – Thunderbird Area



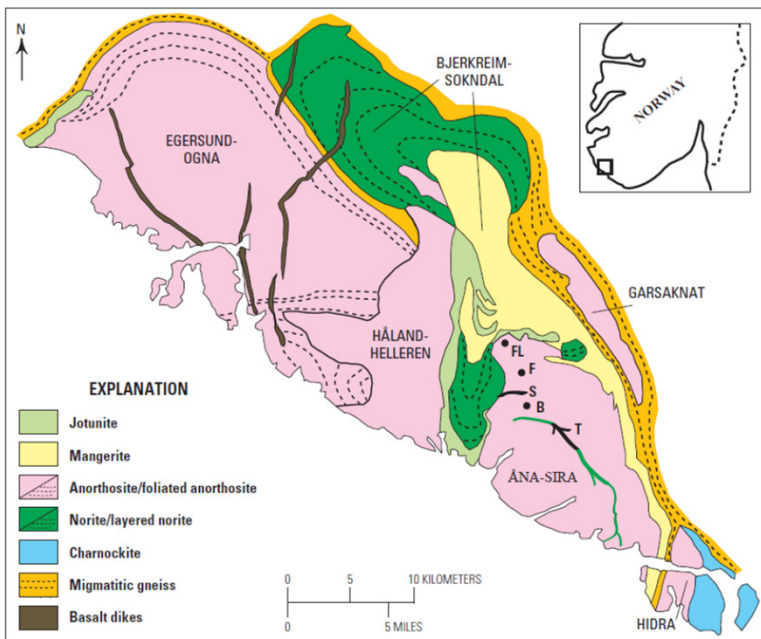
- ❖ **Fe-Ti-V mineralized horizons**
 - ❖ 0.36% V₂O₅ over 178 m including higher grade @ 0.62% V₂O₅ over 16 m
 - ❖ 0.37% V₂O₅ over 100 m including higher grade @ 0.57% V₂O₅ over 21 m
- ❖ Two main mineralized intervals
 - ❖ Structural repetition of the same mineralized interval
- ❖ Fe-Ti-P mineralized horizons outward
- ❖ Presence of a fold axis in the middle of the Thunderbird intrusion
 - ❖ **Based on geochemistry profile, mineral chemistry, and assays profiles**

33



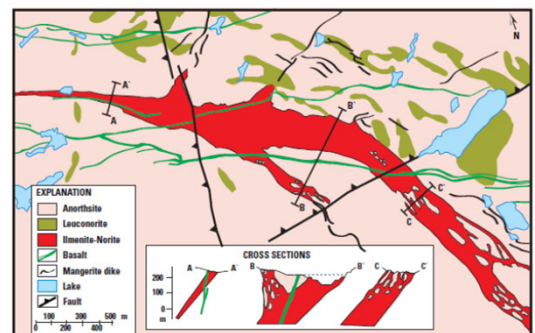
34

Rogaland Anorthosite



❖ Rogaland Anorthosite

- ❖ Includes 4 Anorthositic Massifs and 1 layered intrusion
- ❖ Contains 5 deposits (AS Massif) and 7 deposits (HH Massif)
- ❖ Dominated by hemo-ilmenite oxides



Iron-titanium-oxide deposits:

T=Tellnes
B=Blåfjell, S=Storgangen,
F=Flordalen, and FI=Frøtlog

Woodruff et al. 2013
Canada

Fe-Ti Deposits within Rogaland Anorthosite

Table 3 Principal characters of the Fe-Ti high-grade deposits in Rogaland

Name	Setting	Rock type	Oxide paragenesis	Suggested origin
Tellnes (T)	Dyke-like intrusion in ÅS anorthosite	Homogeneous ilmenite norite, locally laminated	Hemo-ilmenite ± magnetite	Cumulate enriched by crystal sorting in a noritic liquid
Blåfjell (B) Laksedal (L)	Bodies in noritic pegmatite intruded in ÅS	Massive ilmenite	Hemo-ilmenite (Cr-, V-, Mg-rich) ± magnetite (only in L)	Cumulate in a noritic liquid
Storgangen (S)	Concordantly layered dyke intruded in ÅS	Melanoritic layers in slightly deformed layered norite	Hemo-ilmenite + Cr-, V-rich magnetite	Cumulate in differentiated sill
Kydlandsvatn (Ky)	Strongly dipping ilmenite lenses and layers in the contact zone between EGOG and HH	Ilmenite layers in layered anorthosite and leuconorite	Hemo-ilmenite (medium in Cr, V, Mg) ± magnetite	Cumulate in a differentiated sill plastically deformed by the anorthosite emplacement process
Svånes (Sv)	Deformed layered body intruded in HH	Same	Hemo-ilmenite (medium in Cr, V, Mg)	Same
Rødemyr I (RI)	Dyke intruded in HH close to the contact with the NGZ	Oxide minerals vein	Hemo-ilmenite (low Cr) + magnetite (V-rich)	Probably cumulate
Rødemyr II (RII)	Raft in the NGZ (RII) or at the contact with HH	Oxide-rich lenses	Ilmenite (low Cr) + Ti-magnetite + apatite (medium REE)	Cumulate or immiscibility
Kagnuden (Ka)	Veins intruded in HH	Vein of nelsonite with planar orientation of apatite and oxide mineral	Ilmenite (Cr-poor) + Ti-magnetite + apatite (REE-rich)	Immiscibility of a P-, Ti-, Fe-rich liquid
Hestnes (H) Eigerøy (E)				
Jerneld (J)	Veins in HH	Massive hemo-ilmenite	Hemo-ilmenite (very Cr-, Mg-rich)	High temperature solid state segregation and recrystallisation (dynamic recrystallisation)

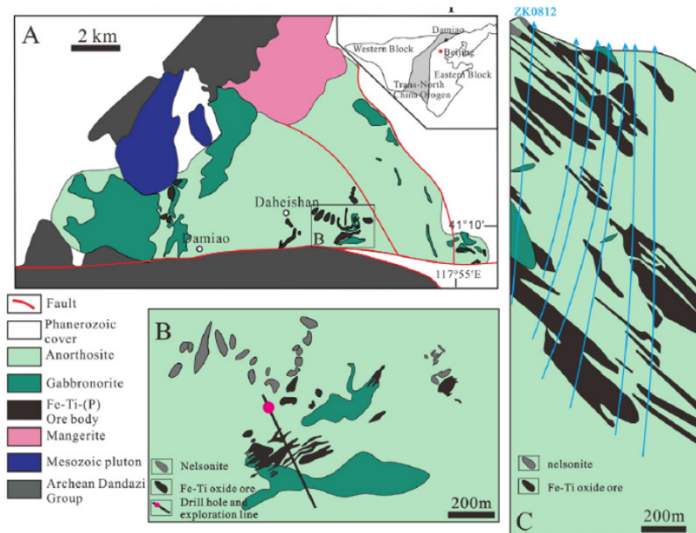


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Duchesne, 1999

Canada

Damaio Anorthosite

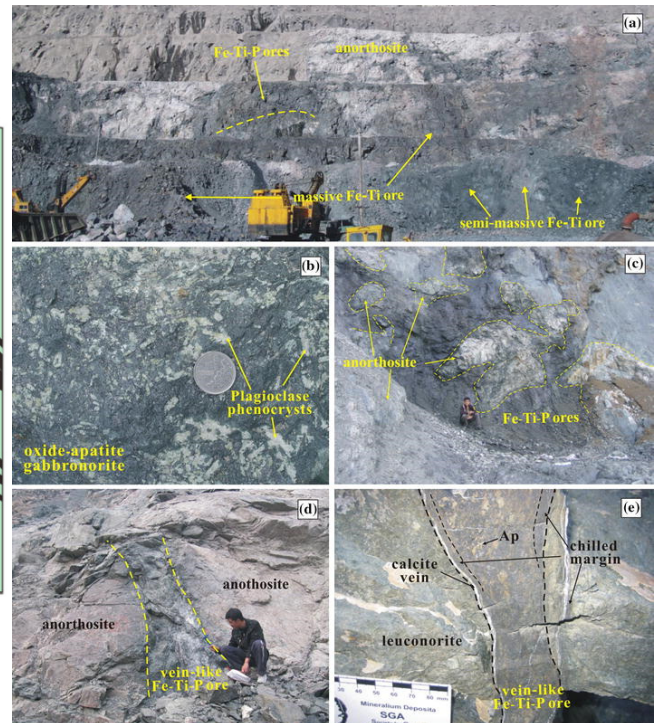


He et al. 2016 - OGR



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Classification Scheme for Fe-Ti-V-(P) Deposits

Major characteristics of magmatic Fe-Ti oxide deposits worldwide.

Association	Host rock	Ore occurrence	Ore mineralogy ^a	Parental magma	Example
Proterozoic anorthosite complexes	Anorthosite, leuconorite, leucogabbro, leucotroctolite	Cross-cutting, massive lenses with sharp contacts against host rocks	Hemo-ilmenite or titanomagnetite, with or without apatite	Ferrodiorite ^b	Tellnes (Norway), Lac Tio (Canada)
Upper parts of large layered complexes	Gabbro, ferrogabbro, ferrodiorite	Conformable, laterally extensive layers with either sharp or gradational contacts against host rocks	Titanomagnetite, with or without apatite	Evolved, Fe-rich tholeiite	Bushveld (South Africa), Windimurra (Australia)
Lower parts of sub-volcanic intrusions associated with flood basalts	Gabbro, ferrogabbro, ultramafic cumulates	Conformable, massive lenses or layers with either sharp or gradational contacts against host rocks	Titanomagnetite, with or without ilmenite	High-Ti ferrobalt	Panzhihua (China), Hongge (China)

^a Only primary ore minerals are listed.

^b The term jotunite has also been used to denote the same type of magma.

Bulk compositions related to magmatic Fe-Ti oxide deposits.

	Fe-rich tholeiite	Ferrodiorite	Picrite	High-Ti basalts		
	(1)	(2)	(3)	(4)	(5)	(6)
SiO ₂	49.79	49.78	44.47	46.89	43.93	44.86
TiO ₂	0.82	3.02	2.35	3.30	3.12	3.61
Al ₂ O ₃	15.82	14.37	11.15	10.13	14.80	11.08
Fe ₂ O ₃	1.18	2.88	1.33	1.37	1.53	1.47
FeO	11.83	13.11	12.03	12.33	13.77	13.23
MnO	0.19	0.23	0.19	0.20	0.20	0.20
MgO	6.14	3.39	14.78	10.96	7.20	8.92
CaO	10.93	7.99	11.38	10.97	10.67	9.55
Na ₂ O	2.97	3.29	1.89	1.44	2.24	2.55
K ₂ O	0.25	1.20	0.08	1.98	1.21	1.22
P ₂ O ₅	0.07	0.74	0.34	0.43	0.33	0.52
Mg# ^b	48	32	69	61	48	55

^aMajor and minor element oxides are in weight percent and are calculated to 100% total on a volatile-free basis.

^bMg# = [molar 100 × Mg/(Mg + Fe²⁺)], assuming 10% of total iron oxide is ferric.



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Ore-Forming Processes - Fe-Ti-V-(P) Deposits

Several mechanisms have been proposed to explain the formation of the Fe-Ti oxide layers. Many of the processes involved are analogous to processes which have been put forward for the genesis of chromitite layers...

- ❖ **Fractional crystallisation, crystal sorting & plagioclase buoyancy**
- ❖ **Liquid immiscibility**
- ❖ *Magma mixing/addition*
- ❖ *Change in physical parameters*
 - ❖ *Oxygen fugacity and Pressure*
- ❖ *Injection of crystal slurries*
- ❖ *Mobilization and Remobilization (mainly invoke for Anorthositic Massif)*
 - ❖ *Filter-press compaction, solid-state remobilization, and hydrothermal remobilization*



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Ore-Forming Processes - Fe-Ti-V-(P) Deposits

Mass-balance problem is less an issue for Fe and V in magnetite than there is for Cr in chromitite layers...

- ❖ First of all, whereas Cr is the major element in chromite, V is simply a **strongly compatible element within the magnetite structure**, and it is the behavior and abundance of Fe that controls magnetite formation
 - ❖ For example, if Fe-rich melts contain on the order of 20 wt.% total FeO (Tegner 1997), the enrichment factor for Fe in oxide layers is only 4
- ❖ Thus, formation of a **1-m-thick layer of oxide** would only require the **suspension of 8 m of silicate crystals**. Since magnetite has a density of over 5 g/cm³, and the grain size of primary magnetite is larger than chromite, magnetite may sink the most rapidly in a melt with crystals of multiple minerals
- ❖ However, mass-balance could be an issue again if we deal with 60-m-thick oxide layers as in the Panzhihua - Will require an overlying column of 2.5 km of magma (Yao et al. 2001)



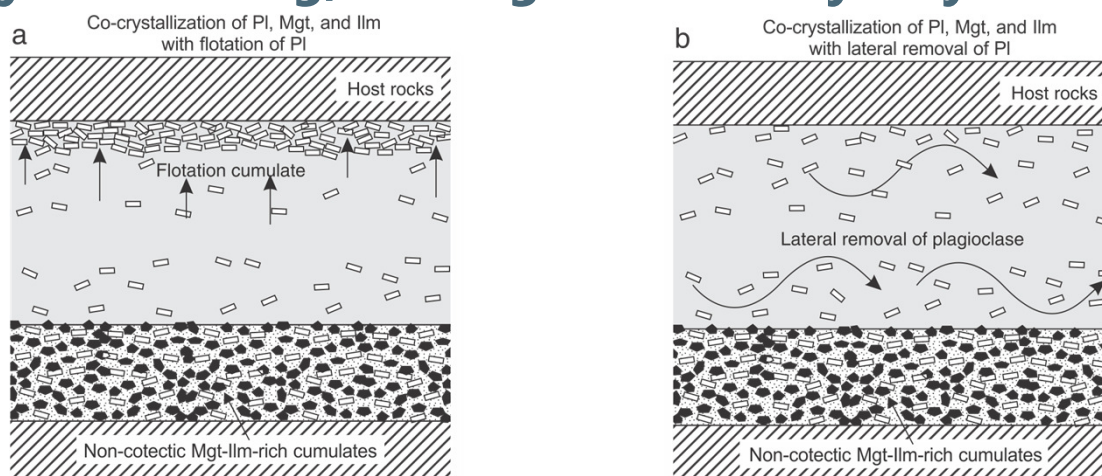
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Cawthorn et al. 2005 – Econ Geol 100th Ann. Vol

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Crystal Settling, Sorting and Pl Buoyancy



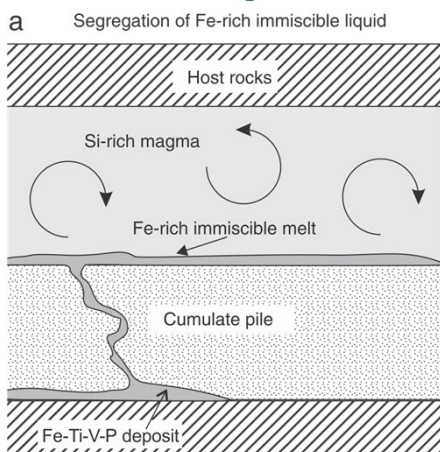
Densities difference after crystallization result in flotation of plagioclase but sinking of the dense Fe-Ti oxides (pilmeneite = 4.7, pmagnetite = 5.2).

Plagioclase can float vertically and form anorthosite that is hard to distinguish from the host anorthosite

Plagioclase be transported laterally to form leucocratic rocks in other parts of the magma chamber

Modified Charlier et al. 2015 - ESR

Immiscibility I.



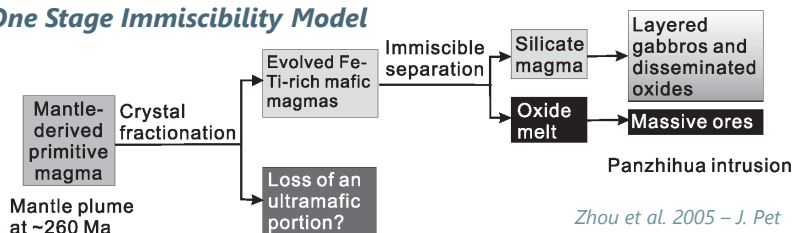
- ❖ Liquid immiscibility is a commonly proposed model for the development of significant massive oxide accumulation
- ❖ Experimental studies using evolved basaltic compositions have produced immiscible liquids, one Fe rich and the other silica rich (Naslund 1983; Philpotts and Doyle 1983)
- ❖ However, in no case did the experimentally produced immiscible liquid contain more than 35 wt.% FeO and still contained considerable normative silicate minerals, not 100% oxide liquid!
- ❖ It has been also argued that if such Fe-oxide liquid were form, it would percolate into interstitial spaces in the underlying cumulate rather than form planar layer with sharp basal contacts (Cawthorn and Ashwal 2009)

- ❖ Alternatively, immiscibility between Fe-rich and Fe-poor silicate liquids has been identified in the Skaergaard and the Sept Iles intrusions based on the presence of contrasting Fe-rich and Si-rich melt inclusions in apatite and olivine in the upper parts of their stratigraphy (Jakobseen et al. 2005, Charlier et al. 2011, Namur et al. 2012)

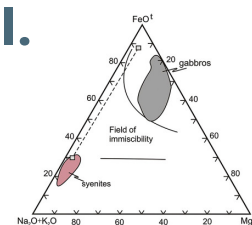
After Charlier et al. 2015; Cawthorn et al. 2005

Genesis of Fe-Ti-V - Panzhihua intrusion II.

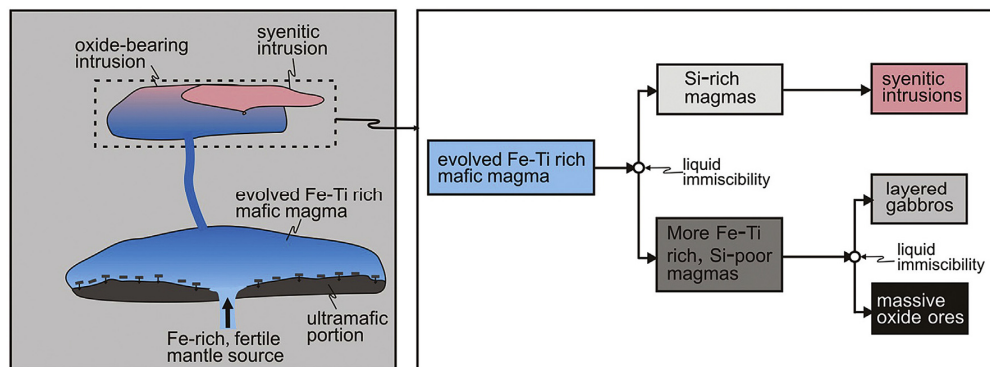
One Stage Immiscibility Model



Zhou et al. 2005 – J. Pet



Two Stages Immiscibility Model



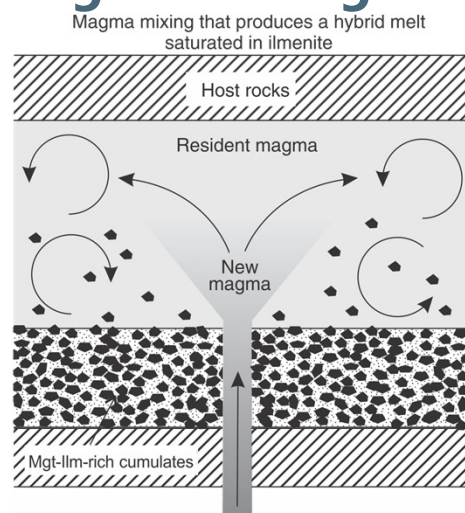
Zhou et al. 2013
– Geo Frontier

Ore-Forming Processes - Fe-Ti-V-(P) Deposits

Change in physical parameters such as Oxygen fugacity and Pressure

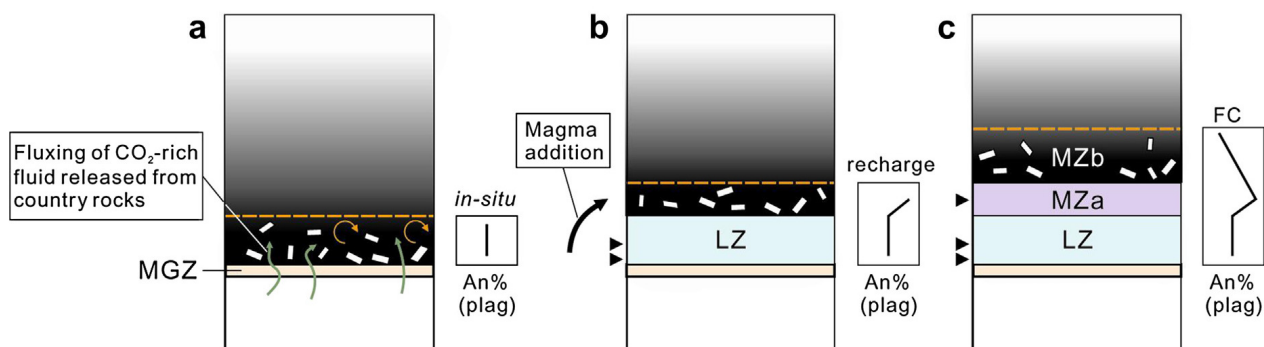
- ❖ An increase in oxygen fugacity (Ulmer, 1969) or total pressure (Cawthorn and McCarthy, 1980) will cause an increase in the magnetite stability field relative to silicate phases, but the source of oxygen, or the driving force for such pressure changes, remains to be identified.
- ❖ Oxidized conditions promoting magnetite crystallization
- ❖ Sources for the introduction of fO_2 into the system have been proposed including CO_2 degassing of footwall rocks (Howarth et al. 2013) and volatile diffusion from underlying cumulates (Klemm 1985)
- ❖ Although a change in fO_2 can account for a change in crystallizing mineralogy but lateral continuity of the magnetite layers necessitates oxygen to be introduced instantaneously and homogeneously, over a long distance...
- ❖ A fluctuation in pressure may result in a uniform pressure change simultaneously across the entire strike length. Several mechanisms to change the pressure including magma addition to the chamber and tectonics have been suggested (Cameron 1978; Cawthorn and McCarthy 1981).

Magma Mixing



- ❖ Introduction of primitive magma may drive the composition of the resultant magma into the spinel stability field where spinel only will crystallize without other minerals thus resulting in monomineralic layer
- ❖ Has only been experimentally shown to be plausible for the crystallization of chromitite layers, this **process may also apply** to magnetite layers **if the addition of new magma will result in magnetite saturation**
- ❖ As example, reversals in plagioclase composition have been documented in the UZ (suggesting magma addition), however, are never coincident with the stratigraphic position of magnetite layers (Ashwal et al. 2005; Tegner et al. 2006; Cawthorn and Ashwal 2009)

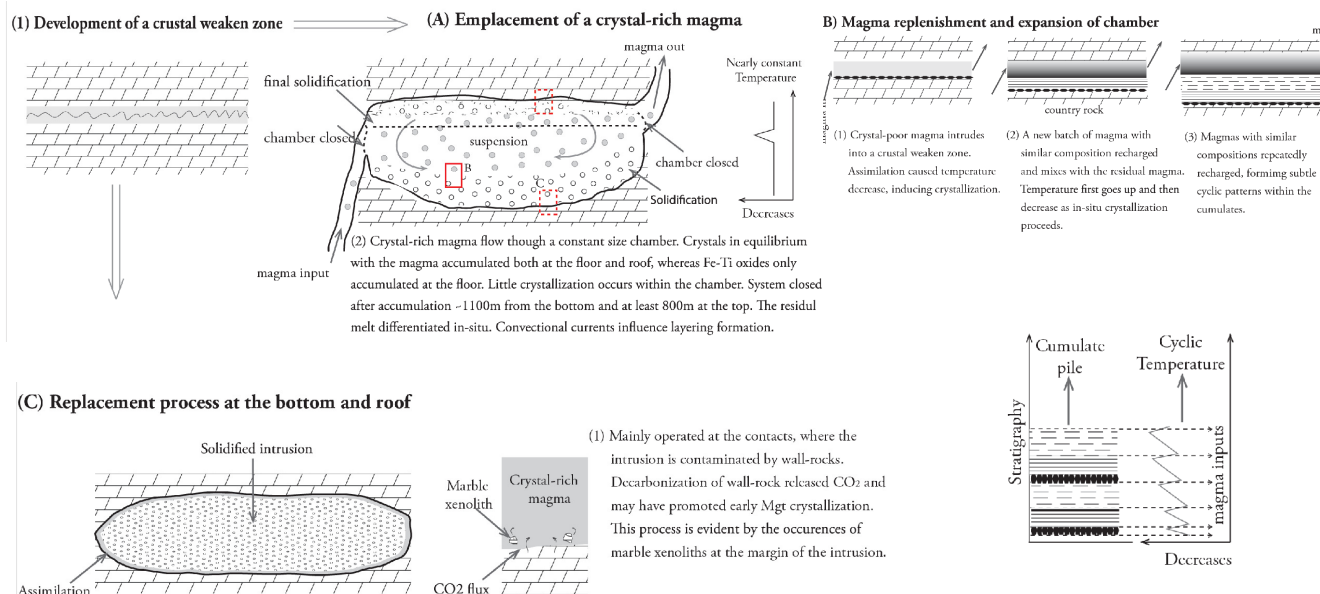
Contamination-Magma Addition-Fractionation



- a) In-situ stage involving fluxing of CO₂-rich fluid from the carbonate country rocks and formation of abundant Fe-Ti oxide ores in a relatively dynamic environment (black triangles).
- b) Recharge stage involving successive addition of magma batches after partial solidification of the LZ and formation of less extensive oxide ore horizons in a relatively static environment; igneous layering is well-developed starting from this stage
- c) Fractionation stage involving largely closed system fractional crystallization and formation of layered gabbros and leucogabbros in the upper MZA and MZb; ore formation is less favorable at this stage and the Fe-Ti oxides are disseminated in the gabbros.

Pang et al. 2013 – Geo Frontier

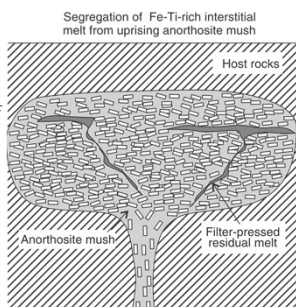
Flow Through System Model



Mobilization and Remobilization

Filter-press compaction

Anorthosite are emplaced as diapirs and crystallize during the ascent of the plagioclase mush. This dynamic emplacement provides the conditions necessary to produce stress-driven melt segregation in partially molten rocks.



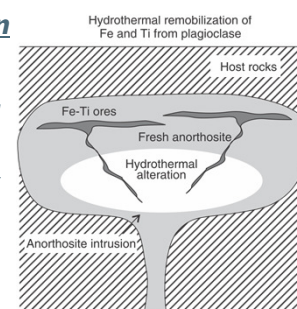
Fe-Ti-P-rich rocks formed by segregation of Fe-Ti-enriched residual melts from uprising pure anorthosite mush (commonly >95% plagioclase).

Solid-state remobilization

It is not rare to observe veins or seams of pure hemo-ilmenite in massif-type anorthosites. However, their emplacement as a melt is highly questionable because the melting point of hemo-ilmenite (at least 1365 °C, depends on dissolved hematite in the solid solution) is hotter than typical magmatic temperature. Ilmenite can easily recrystallize and migrate by diffusion in subsolidus conditions. It can be inferred that such a migration process also occurs on a larger scale during high-temperature deformation (Duchesne 1999).

Hydrothermal remobilization

Plagioclase in anorthosite massifs is commonly altered to various degrees and significant volumes of anorthosite can be affected. Li et al. (2014), who show that alteration of plagioclase reduces the Fe and Ti contents of the anorthosite.



They thus interpret the formation of high-grade Fe-Ti-P ores as a hydrothermal process involving P- and F-rich fluids that migrate, alter the anorthosite and transport Fe and Ti.

After Charlier et al. 2015 - ESR

Concluding Remarks for Fe-Ti-V-(P) Mineralization

Magmatic Fe-Ti-V deposits have formed throughout geological time from a more restricted range of magma types

In order for an economic oxide deposit to form, several processes must occur:

- ❖ *An Fe-rich mafic-anorthositic magma must be generated by fractional crystallization of a more primitive mafic magma*
- ❖ *The magma must be generated at or brought to upper crustal levels prior to reaching oxide saturation*
- ❖ *Magma must mix with other magmas or rocks in a way so as to precipitate much greater amounts of Fe-Ti-V oxides than normal*
- ❖ *The oxides must be segregated and concentrated in a form that is suitable for mining (and preserved from erosion and weathering)*